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# Groundwater age as an indicator of nitrate concentration evolution in aquifers affected by agricultural activities

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## Abstract

The western part of the Drava alluvial aquifer system, located in northern Croatia, contains significant amounts of groundwater, which is primarily used for public water supply and irrigation. The groundwater of this system contains high concentrations of nitrate which is why aquifer system is classified as a groundwater body of poor chemical status under the Water Framework Directive (WFD).

We investigated the groundwater age in this aquifer system and compared the nitrate concentrations in groundwater and the nitrogen pressure from agricultural activity with respect to the estimated mean groundwater age. We used a combination of the environmental tracers: chlorofluorocarbons (CFCs: CFC-12, CFC-11, and CFC-113), sulphur hexafluoride ( $\text{SF}_6$ ), tritium ( $^3\text{H}$ ) and noble gases. By applying lumped parameter models, we determined the groundwater age in aquifers at different depths. On comparing the recharge year, historical data on nitrate concentrations in groundwater, and nitrogen pressure from agricultural activity, we found that these elements are closely related. Our investigation was also supported by the results of numerical simulation of the evolution of nitrate concentration in the saturated zone of the aquifer. The decrease in agricultural pressure caused a decrease in the nitrate concentrations in the youngest, shallow, and oxic groundwater. However, the trend of increasing nitrate concentrations in the deeper part of the aquifer can last for many years. Our research supports the thesis that groundwater age is an important criterion for assessing the

27 effectiveness of protection measures taken in groundwater management and implementation of the  
28 WFD and Nitrate Directive.

29 **Keywords:** porous media, groundwater age distribution, nitrogen pressure, agriculture, nitrate  
30 concentration trend, modelling

31

## 32 **1. Introduction**

33 For sustainable groundwater management, it is vital to have a comprehensive understanding of the  
34 complex and diverse processes of groundwater recharge and the processes that occur during  
35 groundwater flow. There is a large number of terms and phrases in literature on groundwater which  
36 refer to the age and lifetime of a groundwater molecule. For instance, the groundwater age represents  
37 the time taken by a water molecule to reach a specific location, for such as an observation well, from  
38 the moment it was recharged into the subsurface system (Kazemi et al., 2006). The groundwater  
39 residence time is another commonly used term and represents the time taken by water particles to  
40 travel from the recharge area to an aquifer discharge area, such as a river or spring (Modica et al.,  
41 1998). Suckow (2014) provides basic definitions such as the idealised age as one that corresponds to  
42 the results of particle tracking and piston flow model of groundwater flow, mean residence time  
43 involving an age distribution, and apparent age. The idealised groundwater age is the time taken by  
44 water to travel from groundwater table to the sampling site (Torgersen et al., 2013, Suckow, 2014).  
45 This definition is suitable for groundwater age dating methods based on dissolved gases (such as  
46  $^3\text{H}/^3\text{He}$ , chlorofluorocarbons (CFCs), and sulphur hexafluoride ( $\text{SF}_6$ )) because it is a measure of the  
47 time elapsed since the water was last in contact with the atmosphere. Age is generally determined by  
48 matching the measured concentration of dissolved gases in the sample with the corresponding input  
49 concentration of the same gases in a given year of recharge. In nature, the occurrence of various  
50 physical and chemical processes affects the concentrations of dissolved substances in the aquifer;  
51 hence, the groundwater age, estimated using concentrations of dissolved substance, is not necessarily  
52 equal to the time of water transport. In principle, the accuracy of a particular groundwater age depends

53 on how these substances are carried by water. The concentrations of all solutes are, to some extent,  
54 influenced by the transport processes. Their concentrations can be affected by chemical and physical  
55 processes during transport, such as degradation, sorption, diffusion, and dispersion. Therefore, it is  
56 important to understand that the age of groundwater determined from the concentrations of these  
57 tracers actually represents the age of the tracers.

58 The concept of residence time is independent of the definition of idealised age (Suckow, 2014). In  
59 natural groundwater systems, the residence time depends on the position of the water molecule in the  
60 catchment area. For example, the residence time in a gravel layer of high hydraulic conductivity will  
61 vary from the residence time in a clay layer with low conductivity. Considering the residence times  
62 in an actual aquifer system, it is clear that the water samples represent a mixture of idealised ages.  
63 The mean for a mathematically defined mixture of different idealised ages can be calculated using  
64 lumped parameter models (LPMs). Each of such models exhibits its own age distribution curve. For  
65 groundwater management, the shape of the age distribution in these models is more significant than  
66 the actual mean value (Suckow, 2014). For some tracers or tracer combinations, such as  $^{14}\text{C}$  and  
67  $^3\text{H}/^3\text{He}$ , the sample age can be derived using a mathematical formula; Suckow (2014) suggests using  
68 the term ‘apparent tracer age’ to denote the age derived from this method.

69 Groundwater age data can help estimate not only the recharge area, flow path and amount of  
70 groundwater that can be used sustainably, but also the lag of groundwater in relation to a pollution  
71 event, and expected future pollution discharged into groundwater bodies (e.g. Visser, 2009).  
72 Groundwater quality generally improves with a time lag from the cessation of pollution input. The  
73 extent of improvement depends on the type of pollution. In the case of groundwater pollution by  
74 nitrates that are relatively stable in an oxic environment, which is characteristic of the study area, the  
75 duration of poor groundwater quality primarily depends on the mean age of groundwater from the  
76 cessation or reduction of nitrate input. Considering this, groundwater age is becoming an important  
77 tool for assessing the effectiveness of protection measures.

78 Nitrate leaching into groundwater is a major source of pollution caused by intensive agriculture.  
79 Studies conducted in many countries have found that the measured nitrate concentrations are closely  
80 related to agricultural practices (Broers & van der Grift, 2004; Lindsey et al., 2003; Cambray et al.,  
81 2005; Stockmarr et al., 2005; Seifert et al., 2008; van Grinsven et al., 2016). Furthermore, it was  
82 reported that the protection measures applied in agricultural production in Denmark have affected  
83 nitrate concentrations in groundwater (Hansen et al., 2012, 2017).

84 The analysis of the effectiveness of protection measures and trends was primarily considered with  
85 regard to the impact of agricultural production on increased nitrate concentrations in groundwater.  
86 This is because increased nitrate concentrations in groundwater are the among the most pressing  
87 environmental concerns in many countries. A review of the status of the environment and the  
88 consequences of the time lag of nitrate concentrations in groundwater in relation to the time of  
89 application of measures and the load from agricultural production in Europe and North America led  
90 to the conclusion that a water protection policy, aimed at reducing or preventing nitrate pollution in  
91 water, must take into account the time lag of groundwater and the transport of solutes through the  
92 unsaturated and saturated zone (Vero et al., 2018). This lag must be quantified in order to establish  
93 realistic deadlines, thresholds, and expectations, and to plan effective water management practices.  
94 In this regard, it is necessary to determine the groundwater age, that is, the groundwater residence  
95 time, of groundwater polluted by nitrates, in which the measured nitrate concentrations can be related  
96 to the historical input of nitrogen into groundwater. Thus, the reversal trend in agricultural pollution,  
97 as required by the European Union Water Framework Directive (WFD), can be demonstrated, or the  
98 effects of implemented protection measures on nitrate reduction in groundwater can be assessed  
99 (Visser et al., 2009).

100 The study area represents intensive agricultural areas: a combination of mostly (85%) crop  
101 production, intensive vegetable production, and livestock breeding (Bubalo et al., 2014).  
102 Additionally, there are many cattle and chicken farms. High-value, intensively managed crops, such  
103 as vegetables and other irrigated agricultural crops, to which large amounts of nitrogen fertiliser are

104 usually applied, can significantly contribute to nitrate contamination of surface and groundwater  
105 (Bubalo et al., 2014). The application of organic fertilisers (manure) additionally contributes to nitrate  
106 leaching. Higher rates of nitrate losses were recorded in manure treatment than in compost and  
107 inorganic N treatments (Basso & Ritchie, 2005; Thomsen, 2005). Groundwater quality and quantity  
108 monitoring in the Varaždin area has been conducted on an ongoing basis for several years. It has been  
109 observed that the groundwater is oxic and has high nitrate concentrations which are associated with  
110 agricultural production and intensive application of fertilisers and manure; nitrate concentrations have  
111 been increasing in groundwater since the 1970s, and a trend reversal has been achieved in the last 15  
112 years.

113 This paper provides for the first time a comprehensive assessment of the dependence of nitrate  
114 concentration in groundwater on agricultural activity in Croatia. We investigated the groundwater age  
115 in this aquifer system and compared the nitrate concentrations in groundwater with the nitrogen  
116 pressure from agricultural activity with respect to the groundwater age. The atmospheric trace gases  
117 CFC-11, CFC-12, CFC-113, and SF<sub>6</sub>, as well as the radioactive isotope of hydrogen, tritium, and its  
118 product, helium-3, were used for the groundwater age estimation. The production of CFCs started in  
119 early 1940s. Its concentrations gradually began to decrease after the adoption of the Montreal Protocol  
120 in the second half of the 1980s. However, the concentration of SF<sub>6</sub> is still increasing. The high <sup>3</sup>H  
121 concentrations in the atmosphere were consequences of thermonuclear tests, which began in 1952;  
122 peak concentrations of <sup>3</sup>H in rainfall were recorded in 1963 and 1964. The data over the last 20 years  
123 suggest a nearly constant mean annual <sup>3</sup>H concentration (Dulinski et al., 2019; Krajcar Bronić et al.,  
124 2020). We employed the LPM to determine the groundwater age in aquifers at different depths. By  
125 comparing the recharge year, historical data on nitrate concentrations in groundwater, and nitrogen  
126 pressure from agricultural activity, we found that they are closely related to each other. These findings  
127 were supported by the results of numerical simulations of the evolution of nitrate concentrations in  
128 the saturated zone of the aquifer.

129

130        **2. Case study**

131    The Varaždin area is located in the valley of the Drava River at an altitude of 170–200 m above sea  
132    level. It is bounded to the north by the Drava River and to the west and south by mountainous areas.  
133    The valley is dominated by arable land and meadows. The largest settlement in the area is the city of  
134    Varaždin. In the wider study area, two hydroelectric power plants are built on the Drava River with  
135    associated inflow channels, accumulation lakes, engine rooms, and drainage canals (Fig. 1).

136    Fig. 1.

137    According to Köppen's classification of climate types, the Varaždin area belongs to the Cfb climate  
138    type (Šegota & Filipčić, 1996). It has a moderately warm, humid climate with warm summers. The  
139    mean annual air temperatures for the periods 1960–1991 and 1971–2000 were 9.9 °C and 10.2 °C,  
140    respectively (Zaninović et al., 2008). The average annual rainfall for the same period was 879.2 and  
141    843.1 mm, respectively.

142    The Varaždin area is composed of Quaternary sediments within which an alluvial aquifer of  
143    intergranular porosity is formed (Figs. 1, 2). The lithological composition of the aquifer is dominated  
144    by gravel and sand with subordinate silt and clay contents. In the westernmost part of the area, the  
145    hydraulic conductivity of the aquifer reaches 300 m/day (Larva, 2008). It gradually decreases  
146    downstream, approximately reaching 170 m/day in the area east of the Bartolovac pumping site (Fig.  
147    1). The thickness of the aquifer increases from west to east (Fig. 2). The aquifer is covered by a thin,  
148    occasionally absent layer of semipermeable deposits, allowing high amounts of precipitation  
149    infiltration, but also increasing vulnerability of groundwater to pollution by a high degree. The aquifer  
150    is unconfined. In hydrogeological terms, a semipermeable interlayer (aquitard) divides the aquifer  
151    into an upper and a lower aquifers. This aquitard of regional importance appears in the vicinity of  
152    Varaždin and extends further downstream. It is composed of silt and clay with a high sand content  
153    and a thickness of a few meters. However, the layer is occasionally absent.

154    Fig. 2.

155 Groundwater recharge in the study area is mainly achieved through rainfall infiltration. The recharge  
156 rates are relatively high and range from 20% to 30% of annual precipitation, depending on the  
157 thickness and permeability of the covering layer and type of land use (Urumović et al., 1981;  
158 Patrčević, 1995; Brkić, 1999). The Drava River is in direct contact with the aquifer and drains  
159 groundwater under natural conditions. The construction of the hydroelectric power plants on the  
160 Drava River altered the natural conditions; consequently, the aquifer is recharged in the vicinity of  
161 the accumulation lakes, while the drainage canals intensively drain the groundwater (Fig. 1).

162

### 163 **3. Data and methods**

164 In our approach, four input variables (hydrogeological data, historical groundwater quality data,  
165 historical pressure from agriculture data, and environmental tracers) were selected, and then evaluated  
166 and processed. The interrelation between individual segments of the research is shown in Fig. 3.

167 Fig. 3.

168

#### 169 *3.1. Groundwater sampling for CFCs and SF<sub>6</sub> analyses*

170 Groundwater sampling was conducted during two periods: 2017–2018 and 2019–2020 (Table 1).  
171 Eight samples were collected from the observation boreholes at three depths (Table 1). All samples  
172 were collected with a submersible pump Grundfos MP1. Prior to sampling, at least three borehole  
173 volumes of water were pumped out with a pumping rate of about 0.4 L/s. After reaching stable  
174 conditions for temperature, electrical conductivity, and pH, groundwater samples were collected in a  
175 ‘low-flow’ regime with a pumping rate of ~1 L/min.

176 Samples for the analyses of CFCs and SF<sub>6</sub> were collected following the methods described by Oster  
177 et al. (1996). Samples were collected in 500-mL glass bottles stored in containers filled with the same  
178 water to prevent contamination. Groundwater samples were collected in 1-L high density  
179 polyethylene bottles for analyses of <sup>3</sup>H, and in copper tubes with a volume of ~ 40 mL for analyses



180 of noble gases, as described by Weiss (1968). After purifying the copper tubes with at least 10  
181 volumes of tubing, the copper tubes were sealed at both ends with clamps.

182 Table 1

183

### 184 *3.2. Groundwater quality data*

185 Groundwater quality monitoring data were collected from the national monitoring database for which  
186 the Hrvatske vode (a legal entity for water management in Croatia) is responsible, and from the  
187 groundwater quality databases in the investigated recharge areas of the pumping sites (Table 2).

188 Table 2

189

### 190 *3.3. Analysis of CFCs and $^3\text{H}/^3\text{He}$*

191 Analyses of CFCs and  $\text{SF}_6$  were performed in Spurenstofflabor (Wachenheim, Germany) using gas  
192 chromatography, following the methods described by Oster et al. (1996). Tritium and noble gases  
193 were analysed in Isotoptech Zrt. (Debrecen, Hungary). The samples were measured using a Helix  
194 SFT and a VG5400 noble gas mass spectrometer and were determined using the method described  
195 by Palcsu et al. (2010).

196

### 197 *3.4. Groundwater age assessment methods*

198 Groundwater age dating using the atmospheric trace gases is based on Henry's law of solubility.  
199 Using this method, the historical date at which a parcel of water was recharged to a groundwater  
200 system can be calculated. A notable assumption is that, at this point, the water samples were at  
201 equilibrium with the gas concentration in the unsaturated zone. The procedure applied for converting  
202 the measured concentrations in groundwater samples, expressed in picomol per liter (pmol/L) or  
203 femtomol per liter (fmol/L), to the atmospheric equivalent concentration (EAC), expressed in parts

204 per billion per volume (pptv), is described in Kazemi et al. (2006) and IAEA (2006). A recharge  
205 elevation of 200 m and a recharge temperature of 10 °C (Table 2), as well as a groundwater salinity  
206 of 500 mg/L were used. The physical and chemical properties of CFCs were taken from Kazemi et  
207 al. (2006), Cook and Herczeg (2000), Bu & Warner (1995), Cook and Solomon (1995), Warner &  
208 Weiss (1985), and for SF<sub>6</sub> from Cosgrove and Walkley (1981). The obtained EAC values were then  
209 compared to the graphs of CFC concentrations in the air in the Northern Hemisphere  
210 ([https://water.usgs.gov/lab/software/air\\_curve/index.htm](https://water.usgs.gov/lab/software/air_curve/index.htm)), and the year when the analysed CFCs  
211 infiltrated into the groundwater via precipitation was determined (groundwater recharge).

212 To perform groundwater dating with the sparingly soluble SF<sub>6</sub> tracer, excess air must be considered  
213 (Kazemi et al., 2006, Chambers et al., 2019). For the majority of groundwater systems, the excess air  
214 is in the range of ~1–3 cm<sup>3</sup>/L at standard temperature and pressure (STP) (Blusemberg & Plummer,  
215 2000). According to the graph which displays the correction factors for excess air as a function of  
216 recharge temperature (Chambers et al., 2019), the excess air correction factor for SF<sub>6</sub> was assumed to  
217 be 0.8, which corresponds to ~2 cm<sup>3</sup>/L at an STP; the CFC concentrations were not corrected. The  
218 advantages and disadvantages of groundwater dating with CFCs and SF<sub>6</sub> are described in Goody et  
219 al. (2006), Kazemi et al. (2006), and Chambers et al. (2019). Busenberg and Plummer (2000)  
220 considered that the SF<sub>6</sub> method is useful for dating very young groundwater and recharge in urban  
221 environments where CFCs can be elevated owing to local anthropogenic sources.

222 In groundwater younger than the mid-1960s, the highest concentrations of <sup>3</sup>H can no longer be  
223 registered due its radioactive decay into <sup>3</sup>He. However, the apparent age of groundwater can be  
224 calculated from the <sup>3</sup>H/<sup>3</sup>He ratio in a groundwater sample (Schlosser et al., 1988, 1989; Solomon et  
225 al., 1992, 1993; Tolstikhin & Kamenskiy, 1969). <sup>3</sup>H input refers to the <sup>3</sup>H concentration which enters  
226 the saturated zone at the time of recharge. Assuming piston flow conditions, the output represents the  
227 sum of <sup>3</sup>H and <sup>3</sup>He<sub>(trit)</sub> in the sample. The historical <sup>3</sup>H data for precipitation at the Vienna, Ljubljana  
228 and Zagreb Global Network of Isotopes in Precipitation (GNIP) stations were analysed as input data  
229 for <sup>3</sup>H.

230 Groundwater ages were also calculated using LPMs (Maloszewski and Zuber 1996) which are useful  
 231 for describing groundwater systems with a small number of parameters (Maloszewski et al., 1992,  
 232 2002). LPMs are the simplest and most convenient for systems limited to one or two parameters,  
 233 containing young water and concentrations of “modern” tracers such as seasonally variable  $^{18}\text{O}$ ,  $^2\text{H}$ ,  
 234  $^3\text{H}$ ,  $^{85}\text{Kr}$ , CFCs, and  $\text{SF}_6$  as input variables (Maloszewski & Zuber, 1996). In this study, CFC and  $\text{SF}_6$   
 235 data inputs were presented as their atmospheric mixing ratios in precipitation for the Northern  
 236 Hemisphere atmosphere. The historical  $^3\text{H}$  data for precipitation at the Vienna, GNIP station were  
 237 used as input data for  $^3\text{H}$ .

238 The LPMs are represented mathematically as residence time distribution functions or age distribution  
 239 functions  $[g(t)]$  (Maloszewski and Zuber, 1982). A water sample is considered to be composed of  
 240 many parcels with various flow paths leading to the sampling site. Each parcel represents a relatively  
 241 discrete groundwater age and tracer concentration. Mathematically, all LPMs for steady-state flow  
 242 systems with a time-variable tracer input are convolution integrals, as follows:

$$243 \quad C_{out}(t) = \int_{-\infty}^t C_{in}(t') \cdot g(t-t') \cdot e^{-\lambda(t-t')} dt' \quad (1)$$

244 where  $C_{out}$  is the output tracer concentration in groundwater (well, borehole, or spring),  $C_{in}$  is the  
 245 input tracer concentration in the system,  $t$  is the sampling date,  $t'$  is the entry time into the system,  
 246 and  $g(t-t')$  is the residence time distribution or the age distribution function. The term  $e^{-\lambda(t-t')} dt$  is  
 247 related to radioactive decay. The terms  $g(t-t')$  and  $e^{-\lambda(t-t')} dt$  are functions of the idealised age, which  
 248 is, the time difference between infiltration and output time. The equation for the mean age of the  
 249 groundwater sample ( $\tau$ ) is:

$$250 \quad \tau_s = \int_{-\infty}^t (t-t') g(t-t') dt \quad (2)$$

251 In this study, TracerLPM software (Jurgens et al., 2012) was used for the model calculations. Flow  
 252 model calculations were conducted using the piston-flow model (PFM), exponential mixing model  
 253 (EMM), partial exponential model (PEM), and dispersion model (DM) (Maloszewski & Zuber,  
 254 2000). The PFM can be applied to hydrogeologic settings where the flow lines are assumed to have

255 the same residence time, and dispersion and diffusion are negligible. In the EMM approximation, the  
256 flow lines are assumed to have an exponential distribution with regard to the residence time. It is  
257 applicable to - unconfined aquifers of constant thickness, receiving uniform recharge. The PEM is  
258 applicable to the same types of aquifers as for the EMM; however, it is used when only the lower part  
259 of the aquifer is sampled. The DM includes dispersion and advective flow, and can give an  
260 approximate description of age distributions in samples from many aquifer types. It uses the  
261 dispersion parameter (DP), which is the inverse of the Peclet number (Maloszewski and Zuber, 1982,  
262 2002) or the ratio of the dispersion coefficient (D) to the product of velocity (v) and outlet position  
263 (x) (Jurgens et al., 2012).

264

### 265 *3.5. Agricultural N pressure and comparison with nitrate concentrations in groundwater*

266 One of the primary sources of nitrogen, and simultaneously an indicator of the intensity of agriculture  
267 in Croatia, is the fertiliser application per unit area (Romić et al., 2014). The production of mineral  
268 fertilisers, demand and prices of agricultural products, and other specific global circumstances  
269 influenced the increase in the use of plant nutrients through fertilisation, especially after the Second  
270 World War. In Croatia as well in some other EU Member States (e.g., Germany), mineral fertiliser  
271 sales statistics are available at the national level (Ondrašek et al., 2021). A fertiliser factory  
272 Petrokemija d.d., established in 1968, is a major source of historical data on fertiliser application. As  
273 for manure, poultry farms are the primary source in the study area. This study uses available historical  
274 data on the production and fertiliser application and manure in both Europe and Croatia. The sources  
275 of data were expert reports (Romić et al., 2014), published papers (van Grinsven et al., 2015; Dalgaard  
276 et al., 2014; Hansen et al., 2011, 2012, 2017), and statistical data available at Food and Agriculture  
277 Organization (FAO) (<http://www.fao.org/faostat/en/#data/EF>), and Croatian Bureau of Statistics  
278 (CBS) ([https://www.dzs.hr/hrv/system/stat\\_databases.htm](https://www.dzs.hr/hrv/system/stat_databases.htm)).

279 The methodology of comparing nitrate concentrations in groundwater with agricultural N pressures  
280 involved several steps. Firstly, collected historical data on nitrate concentrations at each site were

281 reduced to the mean annual concentrations. Then, the sampling year was converted to the  
282 groundwater recharge year assuming a constant groundwater age at each groundwater sampling point  
283 (Hansen et al., 2017). This means that the recharge year (i.e. the year when nitrates entered  
284 groundwater) was calculated for each annual nitrate concentration as the difference between the  
285 sampling year (i.e. the year with the known nitrate concentration at the considered location) and the  
286 groundwater age. Finally, the obtained recharge years with the corresponding nitrate concentrations  
287 in the groundwater were compared with the historical data of agricultural N pressures.

288

### 289 *3.6. Numerical model of groundwater flow and nitrate transport in saturated zone*

290 The evolution of nitrate concentrations in the saturated zone of the aquifer was simulated using a  
291 mathematical model. The governing equations for groundwater flow (5) and solute transport (6) were  
292 solved numerically using the finite-difference method.

293 The three-dimensional transient movement of groundwater of constant density through a porous  
294 material is described as (Krešić, 2007):

$$295 \quad \frac{\partial}{\partial x} \left( K_{xx} \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left( K_{yy} \frac{\partial h}{\partial y} \right) + \frac{\partial}{\partial z} \left( K_{zz} \frac{\partial h}{\partial z} \right) - W = S_s \frac{\partial h}{\partial t} \quad (3)$$

296 where  $K_{xx}$ ,  $K_{yy}$ , and  $K_{zz}$  are the hydraulic conductivities along the x-, y-, and z- axes ( $LT^{-1}$ ), which  
297 are assumed to be parallel to the major axes of hydraulic conductivity,  $h$  is the hydraulic head (L),  $W$   
298 is the volumetric flux per unit volume representing sources and sinks ( $T^{-1}$ ),  $S_s$  is the specific storage  
299 of the porous media ( $L^{-1}$ ), and  $t$  is time (T).

300 In general,  $S_s$ ,  $K_{xx}$ ,  $K_{yy}$ , and  $K_{zz}$  are functions of space, and  $q$  is a function of space and time. Equation  
301 5 describes the groundwater flow under non-equilibrium conditions in a heterogeneous and  
302 anisotropic medium. In the steady-state model,  $\frac{\partial h}{\partial t}$  in the governing equation (5) is zero, and the  
303 computed heads and fluxes are constant over time.

304 General equation of solute transport without chemical reactions in three dimensions is (Zheng

305 &Wang, 1999):

$$306 \quad \frac{\partial(nc^k)}{\partial t} = \frac{\partial}{\partial x_i} \left( nD_{ij} \frac{\partial c^k}{\partial x_j} \right) - \frac{\partial}{\partial x_{ij}} (nv_i c^k) + q_s C_s^k \quad (4)$$

307 where  $C^k$  is the dissolved concentration of the species ( $\text{ML}^{-3}$ ),  $n$  is the porosity (dimensionless),  $t$  is  
308 time (T),  $x_{i,j}$  is the distance along the respective Cartesian coordinate axis (L),  $D_{ij}$  is the hydrodynamic  
309 dispersion tensor,  $v_i$  is the seepage or linear pore water velocity, calculated as  $v_i = q_i/n$ ,  $q_i$  is the  
310 volumetric flow rate per unit volume of aquifer representing fluid sources (positive) and sinks  
311 (negative) ( $\text{T}^{-1}$ ), and  $C_s^k$  is the concentration of the source or sink flux for species k ( $\text{ML}^{-3}$ ).

312 A regional three-dimensional numerical groundwater flow model, spanning  $2400 \text{ km}^2$  and originally  
313 developed in the scope of the research activities focused on groundwater balance in the western and  
314 central Drava Valley, was used to assess nitrate concentration evolution in the study area (Fig. 3).

315 The model was set up within the Groundwater Modelling System platform and simulated using the  
316 MODFLOW 2005 code (Harbaugh et al., 2017). Horizontal discretisation of the model domain was  
317 performed using a grid size of  $500 \text{ m} \times 500 \text{ m}$ . A vertical discretisation was obtained based on four  
318 layers representing the covering aquitard, upper aquifer, aquitard, and lower aquifer. Rivers,  
319 accumulation lakes, and drains were implemented into the model as head-dependent boundaries,  
320 whereas distributed aquifer recharge, groundwater abstraction at pumping sites, no-flow boundary  
321 across the bottom of the modelling domain, and the eastern, western, and southern boundaries were  
322 all modelled as specified flux boundaries. Model parameters were initially assigned according to the  
323 results of pumping tests carried out mostly for the purpose of the pumping site development and were  
324 subsequently adjusted during the calibration process (Table S1).

325 The model was calibrated in a steady-state with the observed groundwater heads obtained from the  
326 network of observation wells (Fig. 4). For calibration purposes, a parameter estimation tool, PEST,  
327 was used (Doherty, 2015). In accordance with the parsimony principle (Hill, 2006), the model was  
328 kept as simple as possible, and complexity was added in the process of calibration when required.  
329 The goodness of fit between simulated and observed heads was evaluated using mean absolute  
330 residual (MAR), root mean squared residual (RMS) and normalised root mean squared residual

331 (NRMS).

332 Fig. 4.

333 A three-dimensional numerical model of nitrate transport in the saturated zone of the aquifer was  
334 established for the western and central parts of the Drava River Valley, where oxic conditions prevail  
335 in the aquifer (Fig. 3). Downstream, the hydrochemical conditions change to anoxic environment, as  
336 result of sedimentation, gradually becoming dominant and leading to denitrification processes, which  
337 were outside the scope of this study. The model was simulated using the MT3D-USGS code (Bedekar  
338 et al., 2016), which is an upgrade to the groundwater flow solution from the MODFLOW code and  
339 has the capability to route solutes through dry cells that may occur in the Newton–Raphson  
340 formulation of MODFLOW (Niswonger et al., 2011) applied in groundwater flow simulation. The  
341 simulation does not consider processes that affect the retardation and decomposition of nitrate, but  
342 only the advection-dispersion transport, which is in line with the prevailing hydrochemical conditions  
343 of the modelling domain.

344 Quantification of the three-dimensional dispersion effects on solute transport requires the definition  
345 of dispersivity, which is scale-dependent (Wiedemeier et al., 1998; Aziz et al., 2000) and can be  
346 determined using laboratory methods, inverse modelling, or empirical expressions. Gjetvaj (1990)  
347 investigated the dispersivity in the catchment area of the Varaždin pumping site by monitoring the  
348 migration of NaCl solution in a radial flow toward the well. Considering the results of that study,  
349 which fall within the range of dispersivities recorded on such a scale (Gehlar et al., 1992), the  
350 longitudinal dispersivity of 100, transverse dispersivity of 10, and vertical dispersivity of 1, were used  
351 for nitrate transport modelling.

352 Nitrate inputs were simulated using the Neumann boundary condition for zero inflow rate, and the  
353 Cauchy boundary condition for nitrate fluxes at the boundaries with watercourses and lakes, and from  
354 agricultural land.

355 Initial concentrations of nitrate in groundwater in 2006 were derived from national groundwater

356 quality monitoring datasets. Using linear interpolation method the nitrate concentration values were  
357 then interpolated and extrapolated in the areas without information about groundwater quality.

358 Assessing the accurate fertilisers consumption is a critical point in N balancing, and the most sensitive  
359 factor in estimating regional (i.e., at municipality/county level) N surplus (Ondrašek et al., 2021).  
360 There are no data on nitrate concentrations in the unsaturated zone of the aquifer in the study area. In  
361 the neighbouring Slovenia, Urbanc et al. (2014) determined the nitrate leaching to groundwater from  
362 different agricultural lands (uncultivated land and forest, non-intensive land use, extensive land use,  
363 intensive land use and nitrate on farmland in Table 3) in the Drava river aquifer system. Both research  
364 areas, Croatian and Slovenian, belong to the same river basin and have similar regional  
365 hydrogeological characteristics. In addition, there are no significant differences in the agricultural  
366 practices of the neighbouring countries regarding the application of fertilizers and manure. Hence,  
367 the results of research in Slovenia were used to estimate the nitrate surplus reaching the groundwater  
368 from different agricultural lands in the study area. According to Romić et al. (2014), the agricultural  
369 land in the modelled area includes the following crops: cereals, corn, sugar beet, soybeans, oilseeds,  
370 potatoes, vineyards, meadows, pastures, sunflower, tobacco, vegetables, cabbage, orchards and  
371 mosaics. Mosaics were used to represent zones where various crops occupy small areas. In the study  
372 area, mosaics occupy around 30%, corn 18%, cereals 7%, vegetables (including potatoes and  
373 cabbage) 3.5%, meadows and pastures 8%, and all other crops less than 1% of the total area, with  
374 some of them not present at all. For the purpose of this study, different crops listed by Romić et al.  
375 (2014) were merged into five agricultural land use classes, based on similar amounts of nitrate  
376 leaching according to Urbanc et al. (2014). In total, five classes were identified with nitrate leaching  
377 ranging from 7.8 to 81.3 mg/L NO<sub>3</sub> (Fig. 5; Table 3). A particular class was created to account for  
378 the increased application of manure to agricultural land in the vicinity of numerous farms. Manure  
379 accumulated over time in farm premises is subsequently spread on agricultural land, and the nearest  
380 plots are most often used for this purpose (Romić et al., 2014).

381 Fig. 5.



382 Table 3

383 The transport simulation was performed for a 15-y period from 2006 to 2021, over which a constant  
384 input from the sources of nitrates in groundwater was assumed. The simulation results were used to  
385 validate the model performance and analyse trends of nitrate concentrations.

386

## 387 **4. Results and discussion**

### 388 *4.1. Hydrochemical features and groundwater quality*

389 Groundwater is of the CaMg-HCO<sub>3</sub> hydrochemical type (Larva et al., 2010). This is the primary water  
390 type which is principally derived from the dissolution of carbonate minerals (calcite and dolomite)  
391 that compose the aquifer. The pH of the groundwater ranges from 7 to 7.5. The electroconductivity  
392 (EC) varies from 380 to 790  $\mu$ S/cm and depends on the amount of dissolved solids in water. Higher  
393 EC values are recorded in the shallower parts of the aquifer system.

394 The average oxygen concentrations in groundwater in the upper and lower aquifer are  $> 3$  mg/L, iron  
395 concentrations are  $< 100$   $\mu$ g/L, and nitrates are  $> 10$  mg/L. Because CFCs may be degraded under  
396 anoxic conditions, it is important that the study of groundwater age using these tracers is conducted  
397 in oxic groundwater.

398 The primary indicator of poor groundwater quality is the high concentration of nitrate. The monitoring  
399 of nitrate concentrations has been ongoing for many years, and the historically measured data are  
400 presented in Figures 6 and 7. Figure 6a shows the nitrate concentrations from the pumping wells  
401 (label B), capturing the upper aquifer; additionally, Figure 6b shows the nitrate concentrations from  
402 the observation wells (label P) that also monitor the upper aquifer. Due to high nitrate concentrations,  
403 groundwater abstraction from the upper aquifer at the Varaždin pumping site was terminated in the  
404 early 2000s. Since then, the nitrate concentrations have been monitored only in the observation wells  
405 PDS-5, PDS-6, and PDS-7 (Fig. 6c). In a meantime, the lower aquifer was captured, and nitrate  
406 concentrations were monitored in the pumping well B-11 (Fig. 6d). At the Bartolovec pumping site,

407 nitrate concentrations in groundwater were initially measured only in a composite sample (mixture in  
408 Figure 7a) that contained groundwater from wells B-1 and B-2 (upper aquifer). Since 2000,  
409 measurements have been carried out for each pumping well separately (Figs. 7a and 7b).

410 The highest concentrations of nitrate in groundwater, approximately 140 mg/L in the upper aquifer  
411 in the area of the Varaždin and Bartolovec pumping sites, were observed in the early 1980s during  
412 the construction of the hydroelectric power plant on the Drava River. Grđan et al. (1991) associated  
413 this rise in nitrate concentrations with the filling of the accumulation lake, rise of groundwater levels  
414 in the hinterland; these anthropogenic activities, consequently resulted in the leaching of the  
415 unsaturated zone.

416 After this period, nitrate concentrations at the Varaždin pumping site remained relatively high,  
417 reaching a peak in the mid-1990s (Figs 6 a, b, and c). However, at the Bartolovec pumping site, nitrate  
418 concentrations in groundwater in the upper aquifer gradually decreased (Fig. 7a), and subsequently  
419 began to rise in the pumping wells B-1, B-2, and B-7 after 2010; however, the same was not observed  
420 in the observation boreholes P-2G and P-3G (Fig. 7a). Urumović et al. (1991) suggested that the  
421 initial decrease in nitrate concentrations, visible on the composite sample at the Bartolovec pumping  
422 site, is a consequence of changes in boundary conditions after the construction of the hydroelectric  
423 power plant and the filling of the accumulation lake. Groundwater recharge from the accumulation  
424 lake led to a decrease in nitrate concentrations at the Bartolovec pumping site. Water from the Drava  
425 River and the accumulation lake does not contain increased nitrate concentrations; hence, the inflow  
426 of nitrate-free water into the aquifer results in lower nitrate concentrations in the groundwater.  
427 Depending on the trajectories of water particles, this impact is very pronounced in some locations  
428 (shallow, 8-m-deep observation wells P-2G and P-3G), but not in others (pumping wells B-1, B-2, B-  
429 7).

430 Fig. 6.

431 Nitrate concentrations in the lower aquifer at the Varaždin pumping site have exhibited an increasing  
432 trend since initial monitoring in 2002 (Fig. 6d). At the Bartolovec pumping site, the nitrate

433 concentrations in the lower aquifer have been measured since 1993 and are less than 20 mg/L (Fig.  
434 7b and 7c). In some pumping wells, a slight increase in nitrate concentration has been recorded since  
435 2012 (Fig. 7b).

436 Fig. 7.

437

#### 438 *4.2. Interpretation of the mean residence times (MRTs) using environmental tracers*

439 The measured concentrations of CFCs, SF<sub>6</sub> (without correction for excess air), <sup>3</sup>H, and noble gases  
440 are given in Table 4. The measured values in pmol/L range from 0.42 to 4.1 for CFC-12, from 0.8 to  
441 30 for CFC-11, and from 0.04 to 0.46 for CFC-113. Samples from boreholes PDS-4, PDS-5, PDS-6,  
442 PDS-7, and P-3D show concentrations of CFC-11 and CFC-12 above the equilibrium value (that is,  
443 CFC excess). These values are shown in bold in Tables 4. Excess CFC values are usually found in  
444 urban areas where there are many pressures on groundwater. The piezometers used for analyses are  
445 located in relatively urbanized area, so CFC excess in the shallow part of the aquifer could have been  
446 expected. In the case of excess CFCs, groundwater dating is not possible. The measured SF<sub>6</sub>  
447 concentrations varied from 1.3 fmol/L in the deep part of the aquifer system to 3.3 fmol/L in the  
448 shallow part of the aquifer system.

449 Table 4

450 Relating the EAC values to the CFC and SF<sub>6</sub> concentrations in the air in the Northern Hemisphere  
451 ([https://water.usgs.gov/lab/software/air\\_curve/index.html](https://water.usgs.gov/lab/software/air_curve/index.html)), the recharge timing was obtained (Fig.  
452 S1).

453 The measured <sup>3</sup>H concentrations in the groundwater samples ranged from 1.63 to 5.05 TU (Table 4).  
454 The measured concentrations of the noble gases were expressed in ccSTP/g (cubic centimetres at an  
455 STP dissolved in 1 g liquid water) (Table S2). The sample from the monitoring well BVP-3D is the  
456 only sample that is, based on the <sup>3</sup>H content (1.63 ± 0.06 TU) and according to the classification

457 suggested by Clark and Fritz (1997), classified as a mix of sub-modern and modern waters. In other  
458 samples,  $^3\text{H}$  content is 4-5 TU, which classifies them as modern waters.

459 Groundwater ages using the  $^3\text{H}/^3\text{He}$  method were compared with  $^3\text{H}$  input data at GNIP stations in  
460 Vienna, Ljubljana, and Zagreb (Fig. S2). The GNIP station in Vienna has the longest set of input data.  
461 In the period of existence of  $^3\text{H}$  data at all three stations, the set of Vienna data fits well with the data  
462 of the other two stations. Given the proximity of all three stations this is not unexpected. Therefore,  
463 the input data from the Vienna station were used for further analyses. The  $^3\text{H}+^3\text{He}_{(\text{trit})}$  values of all  
464 analysed samples match well with the input values of the  $^3\text{H}$  Vienna station. The groundwater age  
465 increases with the sampling depth (Fig. S2).

466 However, some of the groundwater samples display discrepancies between groundwater ages  
467 estimated by different tracers. The groundwater ages of samples PDS-5, PDS-6, and PDS-7 using  
468 CFC-113 suggest greater ages compared to that estimated from the  $\text{SF}_6$  and  $^3\text{H}/^3\text{He}$  methods.  
469 Concentrations of CFC-11 and CFC-12 above the equilibrium value were determined on all three  
470 samples. Although the concentration of CFC-113 is not above the equilibrium value, it is still  
471 significantly high, which is why it shows older water than the one that, given the hydrogeological  
472 conditions of the investigated site, can be. The groundwater ages obtained for the PDS-4 and P-3D  
473 samples using CFC-113 and  $^3\text{H}/^3\text{He}$  were very similar ( $\sim 30$  y). For BVP-3P and BVP-3D samples,  
474 groundwater ages determined using all three CFCs and  $^3\text{H}/^3\text{He}$  were similar (24.2–36 y for BVP-3P,  
475 and 50–62.1 y for BVP-3D, respectively). Comparatively, the groundwater ages determined by  $\text{SF}_6$   
476 are approximately half for the BVP-3D and P-3D samples and 3–5 times less for BVP-3P. Lower  $\text{SF}_6$   
477 groundwater ages may result from a slight excess of air that could originate from low levels of  
478 contamination (Busenberg & Plummer, 2000; Zoellmann et al., 2001, Wilske et al., 2020).  
479 MacDonald et al. (2003) found that due to the lower solubility,  $\text{SF}_6$  concentrations were less buffered  
480 against changes due to unintentional air entry during sampling. Accordingly, it could be concluded  
481 that there was either an unintentional air entry during sampling or that the groundwater sample

482 contained excess SF<sub>6</sub> that could have originated from a low level of pollution. Because the bottles of  
483 the samples did not contain bubbles, the second option is deemed more likely.

484 The consistency of the groundwater ages was verified based on the assumption that the vertical  
485 velocity  $v_{vert}$  at the water table surface is simply a function of the recharge rate  $R$  and the porosity  $p$   
486 according to the relation  $v = R/p$  (Mahlknecht et al., 2001). The travel time  $t$  for vertical movement  
487 from the water table to sampling depth  $d$  is  $t = d / v_{vert} = d \times (p/R)$ . The travel time for groundwater  
488 samples was calculated for the recharge rates from 20% to 30% of the mean annual precipitation and  
489 porosity of 0.23 (Table 5).

490 Table 5

491 The resulting groundwater ages generally increase with depth and are quite consistent, according to  
492 most methods. In all cases, mixing and dispersion were neglected, that is, piston-flow conditions were  
493 assumed. The influence of mixing and dispersion on MRT was analysed using LPMs (Figs. S3-S5).

494 Graphical estimation of the mean groundwater age of the samples PDS-5 and PDS-6 from a shallow  
495 part of the aquifer system is displayed in Fig. S3. The results of the PFM fit well with the measured  
496 tracers <sup>3</sup>H and <sup>3</sup>H<sub>0</sub>. All models are consistent with the <sup>3</sup>He<sub>(trit)</sub> and <sup>3</sup>H/<sup>3</sup>H<sub>0</sub> values of the sample. The  
497 DM is consistent with the SF<sub>6</sub> and <sup>3</sup>H/<sup>3</sup>H<sub>0</sub> values of the sample PDS-5 and a DP of 1. The EMM is  
498 consistent with the SF<sub>6</sub> and <sup>3</sup>H/<sup>3</sup>H<sub>0</sub> values of the sample PDS-6. The PFM, PEM, and EMM are  
499 consistent with the SF<sub>6</sub> and <sup>3</sup>He<sub>(trit)</sub> values of both samples. For all models, the mean groundwater  
500 ages of the samples are < 10 y.

501 Graphical estimation of the mean groundwater age of the samples P-3D and BVP-3P is presented in  
502 Fig. S4. A deeper part of the upper aquifer was tapped by these observation wells. The PEM results  
503 agree with the measured tracers <sup>3</sup>H and <sup>3</sup>H<sub>0</sub> as well as the <sup>3</sup>He<sub>(trit)</sub> and <sup>3</sup>H/<sup>3</sup>H<sub>0</sub> in the sample P-3D. It  
504 yielded an optimised mean age of 25 y with a PEM ratio of 0.1. The EMM and DM models are  
505 consistent with the SF<sub>6</sub> and <sup>3</sup>He<sub>(trit)</sub> values of the sample. The MRT is ~20 y. The CFC-113 and <sup>3</sup>H/<sup>3</sup>H<sub>0</sub>  
506 values of the sample P-3D are the nearest to the PFM model results and the MRT corresponds to ~32  
507 y.

508 In the groundwater sample BVP-3P, the PFM agrees well with the measured tracers  $^3\text{H}$  and  $^3\text{H}_0$  as  
509 well as the  $^3\text{He}_{(\text{trit})}$  and  $^3\text{H}/^3\text{H}_0$ . The mean groundwater age is estimated to be ~25 y. The DM models  
510 with a DP of 1 are consistent with the  $\text{SF}_6$  and  $^3\text{He}_{(\text{trit})}$  values of this sample. However, the models  
511 yield an optimised mean age of ~10 y. The CFC-113 and  $^3\text{H}/^3\text{H}_0$  values do not fit with any of the  
512 models. The PFM agrees well with the measured tracers CFC-12 and CFC-113, and the estimated  
513 mean groundwater age is ~35 y.

514 Graphical estimation of the mean groundwater age of the samples PDS-4 and BVP-3D is presented  
515 in Fig. S5. These observation wells tap the lower aquifer. For the  $^3\text{H}$  and  $^3\text{H}_0$  methods, as well as for  
516 the  $\text{SF}_6$  and  $^3\text{He}_{(\text{trit})}$ , the measured tracers of PDS-4 are located between the PFM and the PEM model.  
517 The mean groundwater age is ~25 y. For the  $^3\text{He}_{(\text{trit})}$  and  $^3\text{H}/^3\text{H}_0$  the measured tracer in the sample  
518 PDS-4 is located closer to the PEM than the PFM model with the MRT of ~25 y. The PFM is  
519 consistent with the CFC-113 and  $^3\text{H}/^3\text{H}_0$  values of the sample and the MRT corresponds to 30 y. The  
520 DM yields an optimised mean age of 70 y with DP of 1 (for CFC-12 and CFC-113). The PFM results  
521 of the groundwater sample BVP-3D agrees well with the measured tracers  $^3\text{H}$  and  $^3\text{H}_0$ , as well as the  
522  $^3\text{He}_{(\text{trit})}$  and  $^3\text{H}/^3\text{H}_0$ . The mean groundwater age is ~70 y. The DM with a DP of 1 is consistent with  
523 the  $\text{SF}_6$  and  $^3\text{He}_{(\text{trit})}$  values of this sample and gives the mean groundwater age of ~60 y. For the CFC-  
524 11 and  $^3\text{H}/^3\text{H}_0$  methods, the measured tracer in the sample BVP-3D is located closer to the PFM  
525 model with the MRT of ~80 y. The CFC-113 and  $^3\text{H}/^3\text{H}_0$  values of this sample do not correlate with  
526 any models. Graphical estimation of the mean groundwater age of the sample BVP-2D does not return  
527 any reliable results.

528 A more accurate determination of the mean groundwater age was found using a calculation of the  
529 best-fit mean age (Table 6). The best-fit mean age is consistent with the applicability of individual  
530 models to specific groundwater flow from a recharge area to a measured position in the observation  
531 wells. Observation wells PDS-5 and PDS-6 tap relatively thin, unconfined aquifers of approximately  
532 uniform thickness in the recharge area, receiving uniform recharge, and accordingly, the best-fit mean  
533 ages are given by EMM and DM. In contrast, the samples PDS-4, BVP-2D, BVP-3P, and BVP-3D

534 were collected from greater aquifer depths and their best-fit mean ages are, therefore, given by PEM  
535 and DM.

536 Table 6

537

#### 538 *4.3. Vertical age profiles*

539 Vertical profiles of tracer values in an aquifer system are among the most valuable data for  
540 understanding the flow system (IAEA, 2013). The resulting CFCs, SF<sub>6</sub>, and <sup>3</sup>H/<sup>3</sup>He ages generally  
541 increase with depth, ranging from ~8 y in the upper aquifer to > 50 y in the lower aquifer. In the  
542 deeper part of the upper aquifer, the MRT was estimated to be ~29 y (mean MRT values for BVP-3P  
543 in Table 6). A downward vertical groundwater velocity ranges from ~0.75 to 2 m/year, as estimated  
544 by dividing the depth of the screen by the <sup>3</sup>H/<sup>3</sup>He ages of the samples, ignoring the dispersion (Table  
545 5).

546 The covering aquitard above the aquifer is made of clay, silt, and sand, and its thickness ranges from  
547 0 to 2 m. The thickness of these deposits increases toward the southern boundary of the Drava River  
548 Valley. The thickness of the unsaturated zone in the study area ranges from 2 to 6 m. According to  
549 research conducted in similar hydrogeological conditions upstream of the study area, in Slovenia, it  
550 was found that the mean flow velocity through the unsaturated zone approximately ranges from 0.01  
551 to 0.026 m/day (Koroša et al., 2020). In accordance with the thickness of the unsaturated zone, the  
552 total travel time through the unsaturated zone lasts between 0.2 and 1.6 y. Considering the relatively  
553 short duration of the flow through the unsaturated zone, it can be concluded that its addition does not  
554 significantly increase the groundwater age and is practically within the error of the estimated age of  
555 water using a tracer.

556

#### 557 *4.4. Groundwater age distributions*

558 Groundwater age distributions in the selected samples PDS-5 (the shallowest part, upper aquifer),  
559 BVP-3P (deeper part of the upper aquifer), and BVP-3D (the lower aquifer) were analysed and are  
560 shown in Figures 8, 9, and 10. The age distribution displays the fractional contribution of each sub-  
561 parcel of water that collectively constitute the entire sample (Jurgens et al., 2012). The EMM and DM  
562 yield similar age distributions in the sample PDS-5, and both indicate that water contributing to the  
563 well has a distribution of ages from a few to 70 y (Fig. 8a). The models indicate that the amount of  
564 young water (~8 y) reaching the well is 80% for the DM and 70% for the EMM, and only ~20% of  
565 the water from this well is > 10 y (Fig. 8b).

566 Fig. 8.

567 The PEM and DM yield similar age distributions for the sample BVP-3P. Both models indicate that  
568 water samples show a distribution of ages from a few to ~43 y for the PEM, and several times more  
569 for the DM (Fig. 9a). The amount of water < 30 y was ~75% for all models, and < 10% of the water  
570 in this sample was > 100 y according to the DM (Fig. 9b).

571 Fig. 9.

572 The PEM and DM indicate that water contributing to BVP-3D shows a distribution of ages from a  
573 few years for the DM or ~20 y for the PEM, to > 500 y for the DM and ~150–180 y for the PEM (Fig.  
574 10a). The amount of water younger than 50 y reaching the well was ~ 60% for the DM, and ~30%  
575 for the PEM. Approximately 20% of the water in this sample was older than 100 y for the DM, and  
576 ~ 50% for the PEM (Fig. 10b).

577 Fig. 10.

578

#### 579 *4.5. Agricultural N pressures on groundwater and relationship with nitrate concentration trend*

580 The significant use of mineral fertilisers in Croatia began in the 1960s (Romić et al., 2014). According  
581 to Petrokemija d.d., the consumption of mineral fertilizers since 1995 has varied from 400 to 450  
582 thousand tons per year, which is significantly less than the consumption before 1990 (Fig. 11a).



583 According to the FAO database, these amounts are even lower, but the trend is reversed. This  
584 difference is probably due to the fact that Petrokemija shows the amount of fertiliser sold while the  
585 FAO statistics relies on the consumed fertiliser amounts. Nevertheless, the use of mineral nitrogen  
586 fertilisers in Croatia in the period 2012–2015 was decreased by 30% compared to 2008–2011 (EC,  
587 2018).

588 Fig. 11.

589 The fertilizer application curve for the EU is similar to the curve of historical fertiliser application  
590 according to Petrokemija (Fig. 11a). In both curves, it is clear that in the late 1980s, there was a  
591 relatively sharp decline in fertiliser application. A similar distribution of fertiliser use has also been  
592 observed in Denmark (Hansen et al., 2010, 2012, 2017). Van Grinsven et al. (2015) state that there are  
593 several reasons for the reduction in fertiliser application in the EU. First, after the collapse of the  
594 Soviet Union in 1991, the transition of a centrally planned system of large collective state farms to a  
595 market-based system in Eastern European countries ceased the subsidies for the purchase of  
596 fertilisers. Second, MacSharry's 1992 EU reform (the common agricultural policy) reduced  
597 commodity support for cereal production and introduced mandatory land provisions in Western  
598 European countries, which reduced fertiliser demands (Stouman-Jensen et al., 2011). Third, the  
599 European Union imposed strict regulations on N use in agriculture in the 1970s. Furthermore, during  
600 the Homeland War in Croatia in the 1991–1995 period, agriculture, like all other industries, was  
601 neglected, which resulted in a sharp decline in fertiliser sales and consumption.

602 Van Grinsven et al. (2015) also demonstrated that the time distribution of pressures from mineral  
603 fertilisers and manure in the EU countries display a similar trend, which suggests that the pressure  
604 from manure from the 1990s has been declining. The data provided by the FAO database indicate  
605 similar trends (Fig. 11b). There is no historical data on manure consumption in Croatia, but in  
606 accordance with the already described situation, the same trend is expected, as in the rest of the EU.  
607 The manure consumption data in Croatia collected for the last decade, do not indicate any pronounced  
608 trend of change (Fig. 11a).

609 There are also no long-term historical data on the N surplus in Croatia. These data have been collected  
610 since 2010 and indicate that the N surplus averages at ~65% of the fertiliser application (Fig. 11a).  
611 The N surplus generally follows the trend of N input (Fig. 11b). According to research conducted  
612 from July 2011 to December 2012 in the Varaždin area, a total N content of 150–700 kg/ha was found  
613 in the 1 m soil profile, indicating excessive fertiliser use (Bubalo et al., 2014). A maximum of 32%  
614 leached N suggests that there is a direct agricultural influence on groundwater quality.

615 In accordance with aforementioned, the assessment of nitrate evolution in groundwater was made on  
616 the basis that fertiliser application according to Petrokemija reflects the total agricultural pressure in  
617 Croatia. In this study of the dependence of historical nitrate concentrations in groundwater (Figs. 6c,  
618 6d and 7c) on the mean values of groundwater age (Table 6), agricultural pressure trends were crucial,  
619 and not their absolute amounts.

620 Considering the MRT of 8 y in the upper aquifer at the catchment area of the Varaždin pumping site,  
621 a relatively good agreement of the trends is observed for the historical pressure from agricultural  
622 activity and nitrate concentrations in groundwater (Fig. 12). The trends also coincide in the lower  
623 aquifer tapped at the Varaždin pumping site for the mean value of groundwater age of 24 y.

624 Fig. 12.

625 Grđan et al. (1991) concluded that the abrupt increase in nitrate concentrations in groundwater  
626 determined in 1982 could not be related to the application rate of mineral fertiliser applied because  
627 the consumption of mineral fertilisers in the Varaždin area began to decrease due to poor market  
628 opportunities. The authors corroborate this with the fact that the application rate of mineral fertilisers  
629 applied in the early 1980s was lower than that in the 1970s. Despite higher consumption, then the  
630 nitrate concentration in groundwater was ~66 mg NO<sub>3</sub>/L, and in 1982 was 98 mg NO<sub>3</sub>/L (Fig. 4).  
631 However, the authors did not consider the MRT. Given the high consumption of mineral fertilisers  
632 recorded in the mid-1970s, and the MRT of ~8 y, and with the increase in groundwater levels caused  
633 by the filling of the accumulation lake, an increase in nitrate concentration in the early 1980s is not  
634 surprise. Trends similar to those in Fig. 11a are also observed on comparing the quantities of applied

635 mineral fertiliser listed in Grđan et al. (1991) with the corresponding nitrate concentrations in the  
636 recharge year (Fig. 13).

637 Fig. 13.

638 In the upper aquifer tapped at the Bartolovec pumping site for the mean value of groundwater age of  
639 29 y the trends of historical pressure from agricultural activity from early 70s and nitrate  
640 concentrations in groundwater display relatively good agreement (Fig. 14). The high nitrate  
641 concentration in groundwater registered in the early 1980s (Fig. 7a) cannot be related to the  
642 agricultural pressure curve in Fig. 14. At that time, only shallow pumping wells B-1 and B-2 were in  
643 operation. The sudden increase in groundwater levels is caused by the filling of the accumulation  
644 lake. Shallow water table make nitrate entering groundwater more easily and consequently  
645 determined the nitrate loading from vadose zone (Guo et al., 2006). Subsequent changes of the  
646 boundary conditions and the catchment area of the Bartolovec pumping site, and the inflow of nitrate-  
647 free water, allowed decreasing of the nitrate concentration. However, with a dominant groundwater  
648 inflow of the MRT of 29 years from more remote, primarily western areas of the catchment, the nitrate  
649 concentration in groundwater began to increase again, which corresponds to the considered  
650 agricultural pressure (Fig 14).

651 Fig. 14.

652 The results of the assessment of nitrate evolution are supported by the results of the groundwater age  
653 distribution of in the aquifer system. Groundwater in the shallowest part of the system contains  
654 approximately 80% of water younger than 8 years (Fig. 8), and the mean annual concentration of  
655 nitrate in the water reached 100 mg/L (Fig. 12). In the deeper part of the upper aquifer (Bartolovec  
656 pumping site), groundwater contains approximately 75% of water younger than 30 years (Fig. 9), and  
657 the mean annual concentration of nitrate in the water reached 30 mg/L (Fig. 14). In the lower aquifer,  
658 the groundwater age is determined to be approximately 24 years (Varaždin pumping site) and more  
659 than 50 years (Bartolovec pumping site). The aquitard separating the upper and lower aquifers is not  
660 continuous in the Varaždin area (Fig. 2), and consequently, the concentration of nitrate in

661 groundwater reached 60 mg/L. At the Bartolovec pumping site, groundwater in the lower aquifer  
662 contains up to 60% of water younger than 50 years (Fig. 10), and nitrate concentrations reach 15  
663 mg/L (Fig. 7c). According to the applied methodology, the increase in nitrate concentrations in this  
664 aquifer can last for years.

665 This study demonstrates a clear relationship between changes in agricultural pressure, and change in  
666 the nitrate concentrations with the similar temporal pattern and trend reversals. These findings are  
667 comparable with previous studies conducted in Denmark (Hansen et al., 2011; 2012) and the  
668 Netherlands (Broers & van der Grift, 2004; Visser et al., 2009). The nitrate concentrations in the  
669 Denmark were decreased after recharge year 1980, and in the Netherlands, as well as in Croatia,  
670 decreasing of the nitrate concentration began approximately 10 years later (around 1990). The  
671 recharge years were defined on the same way. The difference is in the applied agricultural N pressure.  
672 Hansen et al. (2012) has applied N surplus which is considered to be the best indicator of the impact  
673 of agriculture on the environment in a given period (European Environmental Agency, 2005) and  
674 which represents the amount of N deposited in the soil that is not used in production system (plants  
675 do not absorb all fertilizers), which is why the environment is endangered (Dalgaard et al., 2011;  
676 Hansen et al., 2012). Unfortunately, there are no historical data on N surplus in Croatia. As  
677 aforementioned, in this study the nitrate concentration trend was compared with the trends of  
678 application of mineral fertilizers and manure. Given that the curve of historical data of the N surplus  
679 generally has the same trend as the total agricultural pressure (Fig. 11), this approach can be  
680 considered acceptable both in this case study and in all other locations where there is no historical  
681 data of the N surplus, but there is data on total pressure.

682

#### 683 *4.6. Numerical simulation of the nitrate concentrations evolution in the aquifer system*

684 The model performance was evaluated using the calibration error statistics (Fig. S6). According to  
685 calibration results (MAR = 0.70 m, RMS= 0.83 m, NRMS = 0.90%), the model performs reasonably  
686 well and forms a solid basis for a solute transport simulation. The validation of the solute transport

687 model was performed by comparing the simulated and observed nitrate concentrations. The evolution  
688 of nitrate concentration in the shallow aquifer in the catchment area of the former Varaždin pumping  
689 site for the period from 2006 to 2021 is shown in Figures 15. This area is characterised by the highest  
690 initial nitrate concentrations in the study area. The simulated values correspond well with  
691 observations, especially for the wells PDS-5 and PDS-6. For the observation well PDS-7, the  
692 simulated values are altogether overestimated, most probably because of the overestimated initial  
693 nitrate concentrations in the vicinity of the well. However, the direction of the simulated curve follows  
694 the direction of the observed concentrations reasonably well. The simulation curves for all wells  
695 indicate a slight declining trend in nitrate concentration towards the end of the simulation, which is  
696 generally in line with the conclusions drawn from the mean groundwater age estimates and historical  
697 agriculture pressure (Fig. 12).

698 Fig. 15.

699 The aquifer system is thicker in the catchment area of the Bartolovec pumping site, compared to the  
700 Varaždin pumping site. Consequently, vertical stratification in terms of both mean groundwater age  
701 and groundwater quality is more pronounced. The effects of local influences (groundwater  
702 abstraction, seepage from accumulation lake, drains, application of fertilisers, and manure) and the  
703 effects of regional flow, are also reflected in the distribution of nitrate concentrations in the pumping  
704 site and its catchment area, as shown in Figure 5. Thus, the observation wells P-2G and P-3G that tap  
705 shallow parts of the upper aquifer with relatively young groundwater exhibit low nitrate  
706 concentrations. In contrast, groundwater samples from the pumping wells, which are representative  
707 of regional groundwater flow, have distinctly higher nitrate concentrations, even though there are  
708 differences in nitrate concentrations between individual wells as a consequence of the geometry of  
709 their catchment areas.

710 Vertical and horizontal discretisation of the regional three-dimensional numerical model does not  
711 allow for a detailed analysis of the local influences on groundwater flow and nitrate concentration  
712 distribution. In this regard, and in accordance with the aforementioned, observed nitrate

713 concentrations from the pumping wells, which are representative of regional groundwater flow in the  
714 upper and lower aquifers, were used to evaluate the model performance (Fig. 16). The simulated  
715 curve represents the evolution of nitrate concentrations at the abstraction site and is plotted against  
716 nitrate concentrations in individual wells (Fig. 16). It can be seen that the model is unable to reproduce  
717 variations in concentrations over simulation time, primarily because we assumed steady-state  
718 conditions for both groundwater flow and nitrate transport model. However, the simulated values  
719 correspond with the general trend in both the upper and lower aquifers. Both simulation curves  
720 indicate a positive trend of nitrate concentrations towards the end of the simulation, which is  
721 consistent with the trends of observed nitrate concentrations.

722 Fig. 16.

723 The aim of this study was to improve the understanding of temporal changes in nitrate concentration  
724 in groundwater and to investigate the association of nitrate in groundwater with agricultural load. In  
725 the presented approach, the groundwater age dating and historical agricultural pressure data, together  
726 with the results of nitrate transport modelling, were used to interpret the nitrate trends (Fig. 3). This  
727 is the first time that an assessment of the groundwater age was made in the study area, and the findings  
728 were supported by the results of three-dimensional numerical model of nitrate transport through the  
729 saturated zone of aquifer. The results obtained in the study demonstrate that, even in the areas lacking  
730 detailed temporal data on agricultural pressures, the applied methodological approach yields valuable  
731 results that help explain the evolution of nitrate in aquifer system. Groundwater age dating was once  
732 again proved to be an important indicator of the groundwater vulnerability and, together with the  
733 modelling of groundwater flow and solute transport should be an integral part of decision-making  
734 process in the frame of the sustainable groundwater management. The integrated methodological  
735 approach applied can help assess the effectiveness of protection measures and the implementation of  
736 the EU Water Framework Directive and Nitrate Directive.

737

## 738 **5. Conclusion**

739 This study demonstrates that groundwater age provides valuable information for groundwater  
740 management. A relationship was observed between fertiliser application in agriculture and nitrate  
741 concentrations in groundwater with similar temporal trends. Spatial groundwater age causes  
742 variations in the concentration of nitrates. The MRT of groundwater is  $< 10$  y in the shallow part of  
743 the aquifer,  $\sim 30$  y at the monitoring depth of 35 m, and  $> 50$  y in the monitored part of lower aquifer.  
744 Groundwater age distribution in the groundwater samples indicated that the proportion of young  
745 groundwater decreased in the vertical section of the aquifer system. Approximately 80% of the  
746 groundwater at typical monitoring depths of  $\sim 15$  m was younger than 10 y. At a monitoring depth of  
747 35 m,  $\sim 75\%$  of the groundwater was younger than 30 y. In the monitored part of lower aquifer, up to  
748 60% of the groundwater was younger than 50 y.

749 The decrease in fertiliser application caused a decrease in the nitrate concentrations in the youngest  
750 oxic groundwater. However, the trend of increasing nitrate concentrations at the deepest monitoring  
751 site of the aquifer can last for years.

752 The spatiotemporal development of nitrate concentrations in groundwater, simulated by a three-  
753 dimensional numerical model, is consistent with the observations. The results indicate a negative  
754 trend of nitrate concentrations in the shallow aquifer system with young groundwater, and the same  
755 trend is expected in the future based on the simulation curve. This is consistent with the trend of  
756 historic values and conclusions inferred from the mean groundwater age assessment and fertiliser  
757 application relationship. Owing to a coarse grid, in a complex aquifer system the model was unable  
758 to simulate the horizontal and vertical distribution of nitrate concentrations in detail. However, the  
759 simulated curve fits the observations reasonably well and captures the trend of historic values. Future  
760 studies can focus on developing a local numerical model with finer vertical and horizontal  
761 discretisation that will provide a comprehensive comparison of groundwater age distributions derived  
762 from the groundwater flow model with advective particle tracking and environmental tracer analysis,  
763 and facilitate the simulation of future evolution of nitrates in the aquifer system.

764 This study highlights the importance of understanding groundwater age in water resource  
765 management. The implementation of the EU WFD and Nitrate Directive, among several other  
766 directives, often requires various protection measures. However, for effective evaluation of the  
767 implemented measures, it is crucial to keep in mind that the time required for the effects of these  
768 measures to take place depends on the groundwater age and its distribution.

769

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780

## 781 **References**

782 Aziz, C.E., Newell, C.J., Gonzales, J.R., Haas, P., Prabhakar Clement, T. & Yunwei Sun (2000):  
783 BIOCHLOR: Natural Attenuation Decision Support System; User's manual, version 1.0. EPA/600/R-  
784 00/008, U.S. Environmental Protection Agency, Office of Research and Development, 46 p.,  
785 Washington, D.C.

786 Basso, B., Ritchie, J. T. (2005). Impact of compost, manure and inorganic fertilizer on nitrate leaching  
787 and yield for a 6-year maize–alfalfa rotation in Michigan. *Agriculture, Ecosystems and Environment*  
788 108, 329-341.



789 Bedekar, V., Morway, E.D., Langevin, C.D., and Tonkin, M. (2016): MT3D-USGS version 1.0.0:  
790 Groundwater Solute Transport Simulator for MODFLOW: U.S. Geological Survey Software Release,  
791 30 September 2016, <http://dx.doi.org/10.5066/F75T3HKD>

792 Broers, H. P. & van der Grift, B. (2004): Regional monitoring of temporal changes in groundwater  
793 quality. *J. Hydrol.*, 296, 192-220.

794 Bu, X. & Warner, MJ (1995): Solubilities of chlorofluorocarbons 113 in water and seawater. *Deep-*  
795 *Sea Research Part I: Oceanographic Research Papers*, 42/7, 1151-1161.

796 Bubalo. M., Romić, D., Zovko, M. & Kuspilić, N. (2014): Agricultural Impact on Groundwater  
797 Vulnerability to Nitrate in Northern Croatia. *Agric. Conspec. Sci.*, 79 (1), 23-29.

798 Busenberg E. & L.N. Plummer, 2000. Dating young groundwater with sulfur hexafluoride: Natural  
799 and anthropogenic sources of sulfur hexafluoride. *Water Resour. Res.*, v. 36, p. 3011-3030, October  
800 2000.

801 Cambray, A.R., Marks, M.J., Dawes, S., Ward, R., Hallard, M., MacDonald, A.M., Ruddle, O. &  
802 McConvey, P. (2005): Monitoring effectiveness of the EU Nitrate Directive Action Programmes:  
803 Approach by the United Kingdom. RIVM report.

804 Chambers, L.A., Gooddy, D.C. & Binley, A.M. (2019): Use and application of CFC-11, CFC-12,  
805 CFC113 and SF<sub>6</sub> as environmental tracers of groundwater residence time: A review. *Geosci. Front.*,  
806 10, 1643-1652.

807 Clark & Fritz (1997): *Environmental isotopes in hydrogeology*. Lewis Publishers, London, p. 328.

808 Cook, P.G.& Herczeg, A.L. (1998): *Groundwater chemical methods for recharge studies*. Csiro  
809 Publishing. p. 24.

810 Cook, P.G.& Herczeg, A.L. (2000): *Environmental Tracers in Subsurface Hydrology*. Springer, p.  
811 529.

812 Cook, P.G. and Solomon, D.K. (1995). Transport of atmospheric trace gases to the water table:  
813 Implications for groundwater with chlorofluorocarbons and dating krypton 85. *Water Resour. Res.*,  
814 31, 263-270.

815 Cosgrove, B.A. & Walkley, J. (1981): Solubilities of gases in H<sub>2</sub>O and 2H<sub>2</sub>O. *J. Chromatogr. A*, 216,  
816 161-167.

817 Dalgaard, T., Hutchings, N., Dragosits, U., Olesen, J. E., Kjeldsen, C., Drouet, J.,L., & Cellier, P.  
818 (2011): Effects of farm heterogeneity on modelling and upscaling of nitrogen losses and greenhouse  
819 gas emissions in agricultural landscapes, *Environ. Pollut.*, 159, 3183–3192.

820 Dalgaard, T., Hansen, B., Hasler, B., Hertel, O., Hutchings, N.J., Jacobsen, B.H., Jensen, L.S.,  
821 Kronvang, B., Olesen, J.E., Schjørring, J.K., Kristensen, I.S., Graversgaard, M., Termansen, M. &  
822 Vejre, H. (2014): Policies for agricultural nitrogen management—trends, challenges and prospects  
823 for improved efficiency in Denmark. *Environ. Res. Lett.*, 9, 115002.

824 Doherty, J. (2015): Calibration and uncertainty analysis for complex environmental models.  
825 *Watermark Numerical Computing* .

826 Dulinski, M.; Rozanski, K.; Pierchala, A.; Gorczyca, Z.; Marzec, M. Isotopic composition of  
827 precipitation in Poland: a 44-year record. *Acta Geophys.* 2019, 67, 1637-1648.

828 EC (2018): Report from the Commission to the Council and the European Parliament on the  
829 implementation of Council Directive 91/676/EEC concerning the protection of waters against  
830 pollution caused by nitrates from agricultural sources based on Member State reports for the period  
831 2012-2015, (COM(2018) 257 final, 4 May 2018), European Commission, Brussels.

832 European Environmental Agency (2005): Agriculture and environment in EU-15 – the IRENA  
833 indicator report, EEA Report No 6/2005 ISBN 92-9167-779-5, ISSN 1725-9177, EEA,  
834 Copenhagen.

835 Gelhar, L.W., Welty, C. & Rehfeldt, K.R. (1992): A critical review of data on field-scale dispersion  
836 in aquifers. *Water Resour. Res.*, 19(1), 161-180.

837 Gjetvaj, G. (1990): Identification of dispersivity parameters in radial flow. Proceedings of the 10th  
838 Conference of the Yugoslav Society for Hydraulic Research, 436-440, Sarajevo.

839 Grđan D., Durman, P. & Kovačev-Marinčić, B: (1991): Relationship between regime change and  
840 groundwater quality at Varaždin and Bartolovec pumping sites (in Croatian). Geol. vj., 44, 301-308.

841 Guo H., Li, G., Zhang, D., Zhang, X. & Lu. C. (2006): Effects of water table and fertilization  
842 management on nitrogen loading to groundwater. Agric. Water Manag., 82, 86–98.

843 Hansen, B., Thorling, L., Dalgaard, T. & Erlandsen, M. (2011): Trend Reversal of Nitrate in Danish  
844 Groundwater - a Reflection of Agricultural Practices and Nitrogen Surpluses since 1950. Environ.  
845 Sci. Technol., 45(1), 228-34. DOI: 10.1021/es102334u

846 Hansen, B., Dalgaard, T., Thorling, L., Sorensen, B. & Erlandsen, M. (2012): Regional analysis of  
847 groundwater nitrate concentrations and trends in Denmark in regard to agricultural influence.  
848 Biogeosciences, 9, 3277–3286.

849 Hansen, B., Thorling, L., Schullehner, J., Termansen, M. & Dalgaard, T. (2017): Groundwater nitrate  
850 response to sustainable nitrogen management. Sci. Rep., 7:8566. DOI: 10.1038/s41598-017-07147-2

851 Harbaugh, A.W., Langevin, C.D., Hughes, J.D., Niswonger, R.N., and Konikow, L. F. (2017):  
852 MODFLOW-2005 version 1.12.00, the U.S. Geological Survey modular groundwater model: U.S.  
853 Geological Survey Software Release, 03 February 2017, <http://dx.doi.org/10.5066/F7RF5S7G>

854 Hill, M.C. (2006): The practical use of simplicity in developing ground water models. Ground Water,  
855 44, 775-781.

856 IAEA (2006): Use of chlorofluorocarbons in hydrology, p. 277.

857 IAEA (2013): Isotope Methods for Dating Old Groundwater. p. 357.

858 IAEA/WMO: Global network of isotopes in Precipitation. The GNIP database. Available online at  
859 [http://www-naweb.iaea.org/napc/ih/IHS\\_resources\\_gnip.html](http://www-naweb.iaea.org/napc/ih/IHS_resources_gnip.html)

860 Jurgens, B.C., Böhlke, J.K. & Ebarts, S. (2012): TracerLPM (Version 1): An Excel® Workbook for  
861 Interpreting Groundwater Age Distributions from Environmental Tracer Data. U.S. Geological  
862 Survey Techniques and Methods Report 4-F3, p. 60.

863 Kazemi, G., A., Lehr, J.H. & Perrochet, P. (2006): Groundwater Age. John Wiley & Sons, p. 288.

864 Koroša, A., Brenčić, M. & Mali, N. (2020): Estimating the transport parameters of propyphenazone,  
865 caffeine and carbamazepine by means of a tracer experiment in a coarse-gravel unsaturated zone.  
866 *Water Res.*, 175: 1-12. doi: 10.1016/j.watres.2020.115680

867 Kovač, I., Kovačev-Marinčić, B., Novotni-Horčička, N., Mesec, J. & Vugrinec, J. (2017):  
868 Comparative analysis of nitrate concentration in upper and lower aquifer of the Varaždin system of  
869 aquifers (in Croatian). *Radovi Zavoda za znanstveni rad Varaždin*, 28/1, 41-57  
870 doi:10.21857/9e31lh4zem

871 Krajcar Bronić, I., Barešić, J., Borković, D., Sironić, A., Lovrenčić Mikelić, I. & Vreča P. (2020):  
872 Long-term isotope records of precipitation in Zagreb, Croatia. *Water*, 12, 226.  
873 doi:10.3390/w12010226

874 Kreft, A. & Zuber, A. (1978): On the physical meaning of the dispersion equation and its solution for  
875 different initial and boundary conditions: *Chem. Eng. Sci.*, 33/11, 1471–1480. Available online at  
876 URL <http://www.sciencedirect.com/science/article/pii/0009250978851963>

877 Krešić, N. (2007): Hydrogeology and groundwater modelling. CRC Press, p. 807<.

878 Larva, O. (2008) Aquifer vulnerability at catchment area of Varaždin pumping sites. PhD thesis (in  
879 Croatian), University of Zagreb, p. 198.

880 Larva, O., Marković, T., Brkić, Ž. (2010): Groundwater hydrochemistry of the Quaternary alluvial  
881 aquifer in Varaždin region - Croatia // XXXVIII IAH Congress Groundwater Quality Sustainability  
882 Abstract Book / Zuber, A., Kania, J. & Kmiecik, E. (ur.), Krakow, Poland, University of Silesia Press,  
883 135-136.

884 Lindsey, B.D., Philips, S.W., Donnelly, C.A., Speiran, G.K., Plummer, L.N., Bohlke, J.K., Focazio,  
885 M.J., Burton, W.C. & Busenberg, E. (2003): Residence times and nitrate transport in ground water  
886 discharging to streams in the Chesapeake Bay Watershed. USGS, Water-Resources Investigations  
887 Report 03-4035.

888 MacDonald, A.M., Darling, W.G., Ball, D.F. & Oster, H. (2003): Identifying trends in groundwater  
889 quality using residence time indicators: an example from the Permian aquifer of Dumfries, Scotland.  
890 *Hydrogeol. J.*, 11, 504-517.

891 Mahlkecht, J., Schneider, J.F. & Oster, H. (2001): CFCs and SF<sub>6</sub> in aquifers investigations - The  
892 significance of the unsaturated zone and the origin of groundwater. Available online at  
893 [https://www.semanticscholar.org/paper/CFCs-and-SF6-in-aquifers-investigations-The-of-the-](https://www.semanticscholar.org/paper/CFCs-and-SF6-in-aquifers-investigations-The-of-the-Mahlkecht/dbe32e3f8bf0e21529c3654de4786cc3846be18b#paper-header)  
894 [Mahlkecht/dbe32e3f8bf0e21529c3654de4786cc3846be18b#paper-header](https://www.semanticscholar.org/paper/CFCs-and-SF6-in-aquifers-investigations-The-of-the-Mahlkecht/dbe32e3f8bf0e21529c3654de4786cc3846be18b#paper-header)

895 Maloszewski, P. & Zuber, A. (1982): Determining the turnover time of groundwater systems with the  
896 aid of environmental tracers: 1. Models and their applicability. *J. Hydrol.*, 57/3-4, 207-231.

897 Maloszewski, P., Rauert, W., Trimborn, P., Herrmann, A. & Rau, R., (1992): Isotope hydrological  
898 study of mean transit times in an alpine basin (Wimbachtal, Germany). *J. Hydrol.*, 140, 343-360.

899 Maloszewski, P., Zuber, A., 1996. Lumped parameter models for interpretation of environmental  
900 tracer data. In: IAEA (Ed.), *Manual on the mathematical models in isotope hydrogeology*.  
901 International Atomic Energy Agency, Vienna, pp. 9-59.

902 Maloszewski P., Stichler, W., Zuber, A. & Rank, D. (2002): Identifying the flow systems in a karst-  
903 fissured-porous aquifer, the Schneetalpe, Austria, by modelling of environmental <sup>18</sup>O and <sup>2</sup>H  
904 isotopes. *J. Hydrol.*, 256 (1-2), 48-59.

905 Modica, E., Buxton, H.T. & Plummer, L.N. (1998): Evaluating the source and residence times of  
906 groundwater seepage to streams, New Jersey Coastal Plain. *Water Resour. Res.*, 34, 2797-2810.

907 Niswonger, R.G., Panday, Sorab, and Ibaraki, Motomu (2011): MODFLOW-NWT, A Newton  
908 formulation for MODFLOW-2005: U.S. Geological Survey Techniques and Methods 6-A37, 44 p.

909 Ondrašek, G., Bakić Begić, H., Romić, D., Brkić, Ž., Husnjak, S. & Bubalo Kovačić, M. (2021): A  
910 novel LUMNAqSoP approach for prioritising groundwater monitoring stations for implementation  
911 of the Nitrates Directive. *Environ. Sci. Eur.* 33, 23.

912 Oster, H., Sonntag, C., & Münnich, K. O. (1996): Groundwater age dating with chlorofluorocarbons.  
913 *Water Resour. Res.*, 32, 2989–3001.

914 Palcsu L., Major Z., Köllő Z. & Papp L. (2010): Using an ultrapure  $^4\text{He}$  spike in tritium measurements  
915 of environmental water samples by the  $^3\text{He}$ -ingrowth method. *Rapid Commun. Mass Spectrom.*, 24,  
916 698-704.

917 Patrčević, V. (1995) Hydrological analysis of vertical water balance of groundwater at the area of  
918 river alluvium. PhD thesis (in Croatian), University of Zagreb, p. 161.

919 Romić, D., Husnjak, S., Mesić, M., Salajpal, K., Barić, K., Poljak, M., Romić, M., Konjačić, M.,  
920 Vnučec, I., Bakić, H., Bubalo, M., Zovko, M., Matijević, L., Lončarić, Z., Kušan, V., Brkić, Ž. &  
921 Larva, O. (2014): The impact of agriculture on surface and groundwater pollution in the Republic of  
922 Croatia. Faculty of Agriculture, University of Zagreb, p. 302.

923 Schlosser, P., Stute, M., Dorr, H., Sonntag, C. & Munnich, K.O. (1988): Tritium- $^3\text{He}$  dating of  
924 shallow groundwater. *Earth Planet. Sci. Lett.*, 89, 353–362.

925 Schlosser, P., Stute, M., Sonntag, C. & Munnich, K. O. (1989): Tritogenic  $^3\text{He}$  in shallow  
926 groundwater. *Earth Planet. Sci. Lett.*, 94, 245–256.

927 Seifert, D., Sonnenborg, T.O., Scharling, P. & Hinsby, K. (2008): Use of alternative conceptual  
928 models to assess the impact of a buried valley on groundwater vulnerability. *Hydrogeol. J.*, 16, 659-  
929 674.

930 Solomon, D.K., Poreda, R.J., Schiff, S.L. & Cherry, J.A. (1992): Tritium and helium 3 as groundwater  
931 age tracers in the Borden aquifer. *Water Resour. Res.*, 28, 741-755.

932 Solomon, D.K., Schiff, S.L. Poreda, R.J. & Clarke, W.B. (1993): A validation of the  $3\text{H}/3\text{He}$  method  
933 for determining groundwater recharge. *Water Resour. Res.*, 29, 2951-2962.

934 Stockmarr, J., Grant, R. & Jorgensen, U. (2005): Monitoring effectiveness of the EU Nitrate Directive  
935 Action Programmes: Approach by Denmark. RIVM report.

936 Stouman-Jensen, L., van der Hoek, K.W., Polusen, D., Zevenbergen, J.F., Palliere, C., Lammel, J. et  
937 al. (2011): Benefits of nitrogen for food, fibre and industrial production. In: Sutton, M.A. (Eds.),  
938 European nitrogen assessment, Cambridge Univ. Press, Cambridge, UK, p. 612.

939 Suckow, A. (2014): The Age of Groundwater – definitions, models and why we do not need this term.  
940 *Applied Geochemistry*, 50, 222-230.

941 Šegota, T. & Filipčić, A. (1996): Climatology for geographers. Jelić, T. (Ed), Školska knjiga, Zagreb,  
942 471 p.

943 Thomsen, I.K. (2005). Crop N utilization and leaching losses as affected by time and method of  
944 application of farmyard manure. *Eur J Agron.*, 22, 1-9.

945 Tolstikhin, I.N. & Kamenskiy, I.L. (1969): Determination of groundwater ages by the T– $3\text{He}$  method.  
946 *Geochem. Int.*, 6, 810–811.

947 Torgersen, T., Purtschert, R., Phillips, F.M., Plummer, L.N., Sanford, W. & Suckow, A. (2013):  
948 Defining Groundwater Age. In: Suckow, A., Aggarwal, P.K., Araguas-Araguas, L.J. (Eds.), *Isotope*  
949 *Methods for Dating Old Groundwater*. IAEA, Vienna, pp. 21-32.

950 Urbanc, J., Krivic, J., Mali, N., Ferjan, S., T., Koroša, A., Šram, D., Mezga, K., Bizjak, M., Medić,  
951 M., Bole, Z., Lojen, S., Pintar, M., Udovč, A., Glavan, M., Kacjan-Maršič, N., Jamšek, A., Valentar,  
952 V., Zadavec, D., Pušenjak, M., Klemenčič Kosi, S. (2014): Farming opportunities in water protection  
953 areas: final project report. Geological Survey of Slovenia, 154 p., Ljubljana.

954 Urumović, K., Šarin, A. & Donadini, K. (1981): Groundwater balance in the river basin district  
955 Drava-Dunav for 1979-1980 (Technical report – on Croatian). Croatian Geological Survey, Zagreb.

956 Urumović, K., Duić, Ž. & Hlevnjak, B. (2002): Positive and negative influences of hydroelectric  
957 plants on Drava river to groundwater quality. Conference Abstracts, 21st Conference of the Danubian  
958 Countries on the Hydrological Forecasting and Hydrological Bases of Water Management,  
959 Bucharest, Romania, p. 91.

960 van Grinsven, H.J.M., Bouwman, L., Cassman, K.G. & van Es, H.M., McCrackin, M.L. & Beusen,  
961 A.H.W. (2015): Losses of Ammonia and Nitrate from Agriculture and Their Effect on Nitrogen  
962 Recovery in the European Union and the United States between 1900 and 2050. *J. Environ. Qual.*,  
963 44/2, 356-367. <https://doi.org/10.2134/jeq2014.03.0102>

964 van Grinsven, H.J.M., Tiktak, A. & Rougoor, C.W. (2016): Evaluation of the Dutch implementation  
965 of the nitrates directive, the Water framework directive and the national emission ceilings directive.  
966 *NJAS-Wagen J Life Sc.*, 78, 69–84.

967 Vero, S.E., Basu, N.B., van Meter, K., Richards, K.G., Mellanderl, P.E., Healy, M.G. & Fenton, O.  
968 (2018): Review: the environmental status and implications of the nitrate time lag in Europe and North  
969 America. *Hydrogeol. J.*, 26, 7- 22.

970 Visser, A. (2009) Trends in groundwater quality in relation to groundwater age, PhD thesis,  
971 Netherlands Geographical Studies 384, Faculty of Geosciences, Utrecht University, Netherlands.

972 Warner, M.J. & Weiss, R.F. (1985): Solubilities of chlorofluorocarbons 11 and 12 in water and  
973 seawater. *Deep Sea Res. Part A Oceanogr. Res. Pap.* 32/12, 1485-1497.

974 Weiss, R.F. (1968): Piggyback sampler for dissolved gas studies on sealed water samples. *Deep-Sea*  
975 *Res.*, 15, 695-699.

976 Wiedemeier, T.H., Swanson, M.A., Moutoux, D.E. Gordon, E.K., Wilson, J.T., Wilson, B.H.,  
977 Kampbell, D.H., Haas, P.E., Miller, R.N., Hansen, J.E. & Chapelle, F.H. (1998): Technical protocol  
978 for evaluating natural attenuation of chlorinated solvents in ground water. EPA/600/R-98/128, U.S.  
979 Environmental Protection Agency, Office of Research and Development, Washington, D.C.



980 Wilske, C., Suckow, A., Mallast, U, Meier, C, Merchel, S., Merkel, B., Pavetich, S, Rödiger, T.,  
981 Rugel, G., Sachse, A., Weise, S.M. & Siebert, C. (2020): A multi-environmental tracer study to  
982 determine groundwater residence times and recharge in a structurally complex multi-aquifer system.  
983 Hydrol. Earth Syst. Sci., 24, 249–267.

984 Wilson, G.B. & McNeill, G.W. (1997): Noble gas temperatures and excess air component. Appl.  
985 Geochem., 12, 747-762.

986 Zaninović, K.; Gajić-Čapka, M.; Perčec Tadić, M.; Vučetić, M.; Milković, J.; Bajić, A.; Cindrić, K.;  
987 Cvitan, L.; Katušin, Z.; Kaučić, D.; Likso, T.; Lončar, E.; Lončar, Ž.; Mihajlović, D.; Pandžić, K.;  
988 Patarčić, M.; Srnec, L.; Vučetić, V. (2008): Klimatski atlas Hrvatske/Climate atlas of Croatia 1961–  
989 1990, 1971–2000. State Hydrometeorological Institute, Zagreb, Croatia.  
990 [https://klima.hr/razno/publikacije/klimatski\\_atlas\\_hrvatske.pdf](https://klima.hr/razno/publikacije/klimatski_atlas_hrvatske.pdf)

991 Zheng, C. & Wang, P.P. (1999): MT3DMS: a modular three-dimensional multispecies model for  
992 simulation of advection, dispersion and chemical reactions of contaminants in groundwater systems;  
993 Documentation and User Guide, Contract Report SERDP-99-1, U.S. Army Engineer Research and  
994 Development Center, Vicksburg, Mississippi.

995 Zoellmann, K., Kinzelbach, W. & Fuld, C. (2001): Environmental tracer transport (3H and SF6) in  
996 the saturated and unsaturated zones and its use in nitrate pollution management. J. Hydrol., 240, 187–  
997 205.

## Figure captions

Fig. 1. Study area with hydraulic head equipotential lines and groundwater sampling locations (according to Larva, 2008)

Fig. 2. Lithological cross-section of the aquifer system (according to Larva, 2008)

Fig. 3. Flowchart of the study

Fig. 4. Simulated groundwater heads for average hydrological conditions and spatial distribution of initial concentrations of nitrates in groundwater in the upper aquifer

Fig. 5. Spatial distribution of modelled nitrate concentrations in infiltrating water

Fig. 6. Nitrate concentration in groundwater in the catchment area of the Varaždin pumping site: (a) pumping wells with screen in the upper aquifer, (b) observation wells with screen in the upper aquifer, (c) mean annual nitrate concentrations (the upper aquifer), and (d) mean annual nitrate concentrations (the lower aquifer).

Fig. 7. Nitrate concentration in groundwater in the catchment area of the Bartolovec pumping site: (a) pumping and observation wells with screen in the upper aquifer, (b) pumping and observation wells with screen in the lower aquifer, (c) mean annual nitrate concentrations in both aquifers.

Fig. 8. Modelled groundwater age distribution in the sample PDS-5: (a) fraction of sample and (b) cumulative frequency

Fig. 9. Modelled groundwater age distribution in the sample BVP-3P: (a) fraction of sample and (b) cumulative frequency

Fig. 10. Modelled groundwater age distribution in the sample BVP-3D: (a) fraction of sample and (b) cumulative frequency

Fig. 11. Input data on agricultural N: (a) consumption of fertilizers in the EU and Croatia and (b) input data on N fertilizer, N manure and N surplus in EU (<http://www.fao.org/faostat/en/#data/EF>).

Fig. 12. Time series of the mineral fertilizer consumption according to Romić et al. (2014) and mean annual nitrate concentrations in groundwater in the wider area of the Varaždin pumping site

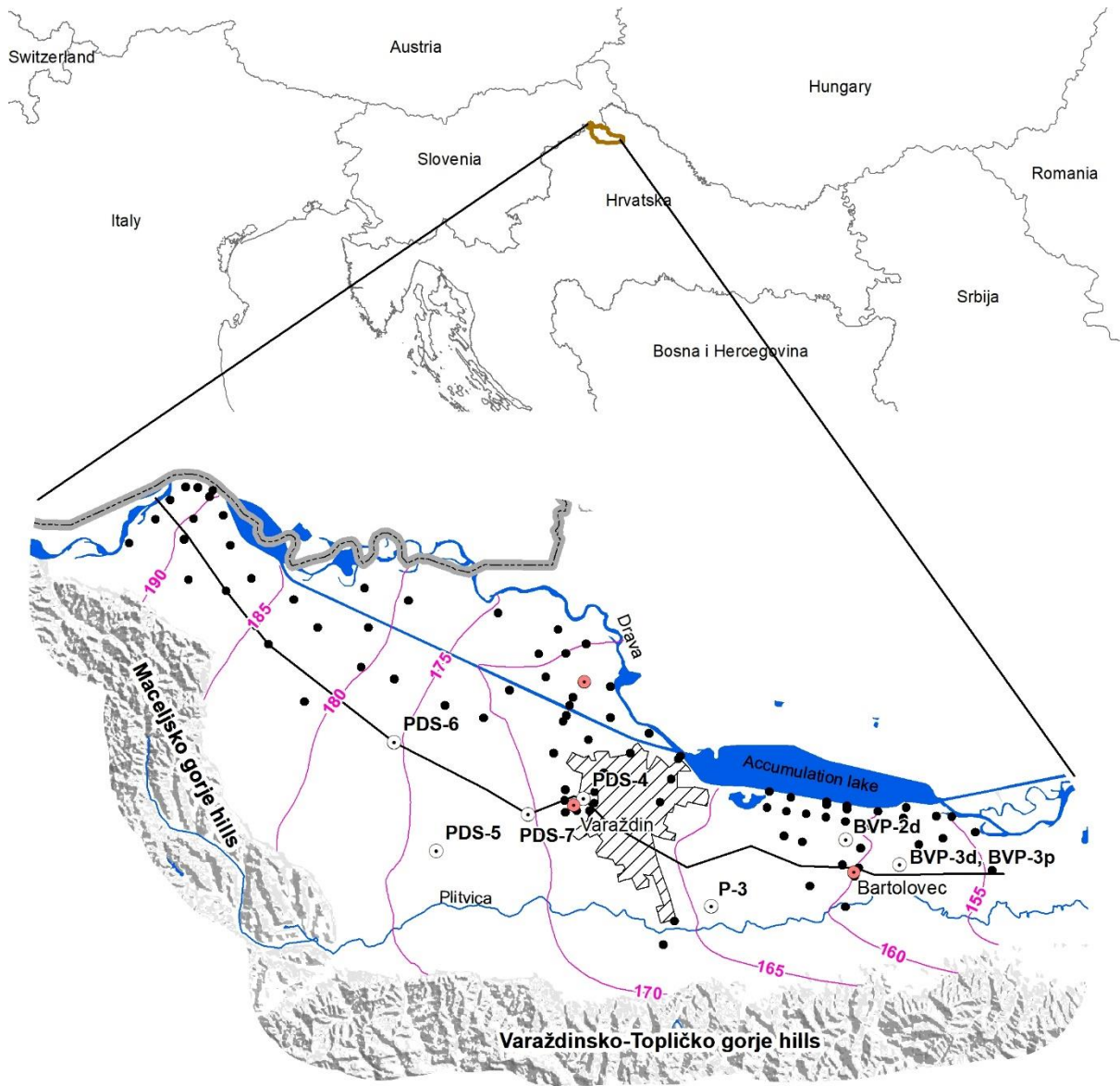
Fig. 13. Time series of the mineral fertilizer consumption in the Varaždin area according to Grđan et al. (1991) and mean annual nitrate concentration in groundwater in the area of Varaždin pumping site

Fig. 14. Time series of the mineral fertilizer consumption according to Romić et al. (2014) and mean annual nitrate concentrations in groundwater at the Bartolovec pumping site

Fig. 15. Simulated and observed nitrate concentrations at observation wells PDS-5 (a), PDS-6 (b) and PDS-7 (c) at the Varaždin pumping site

Fig. 16. Simulated and observed nitrate concentrations in the upper (a) and lower aquifer (b) at the Bartolovec pumping site

Figure 1



### Legend

- pumping site
- groundwater sampling site
- observation well
- urban area
- cross section
- groundwater heads
- state border

Figure 2

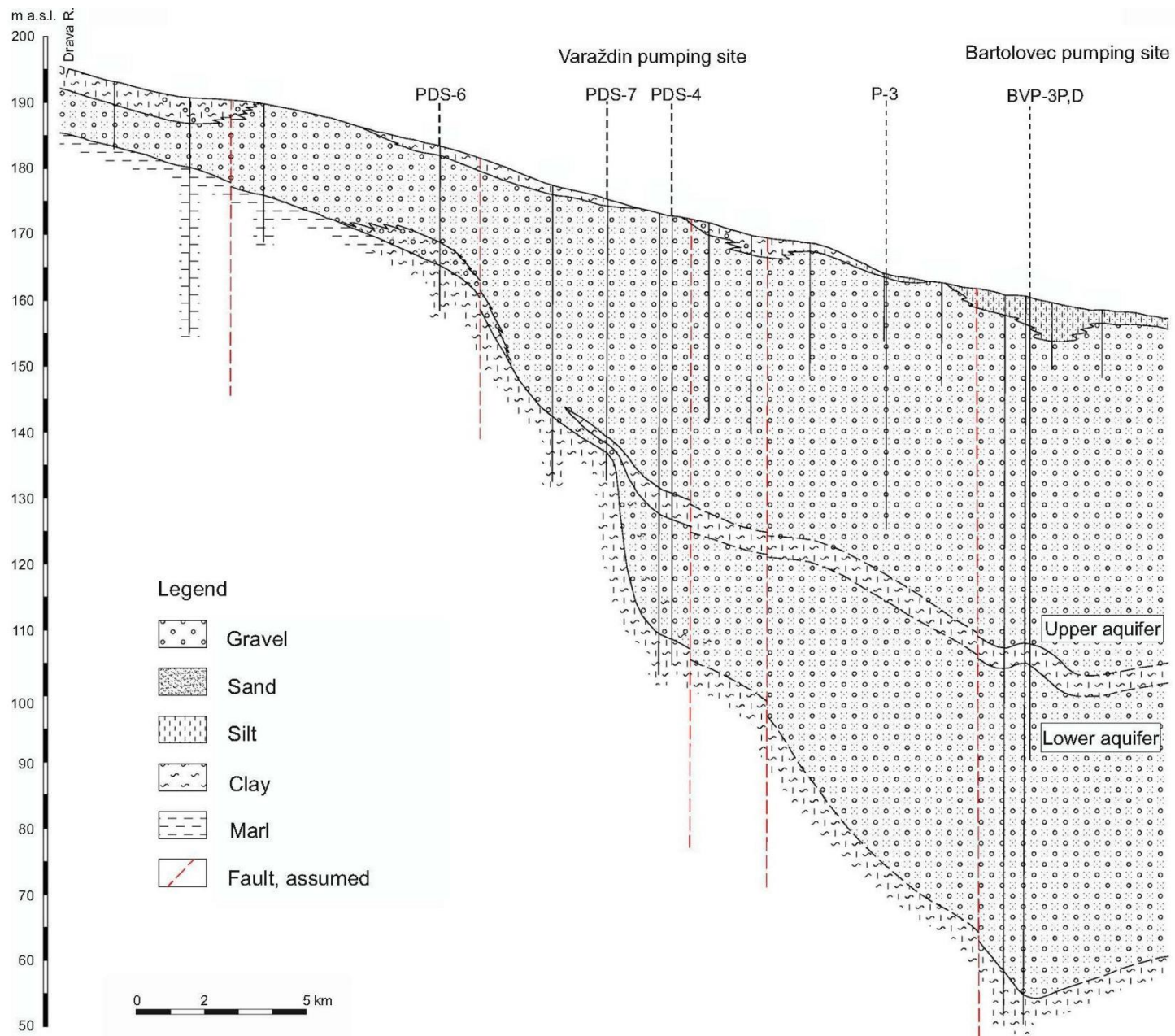


Figure 3

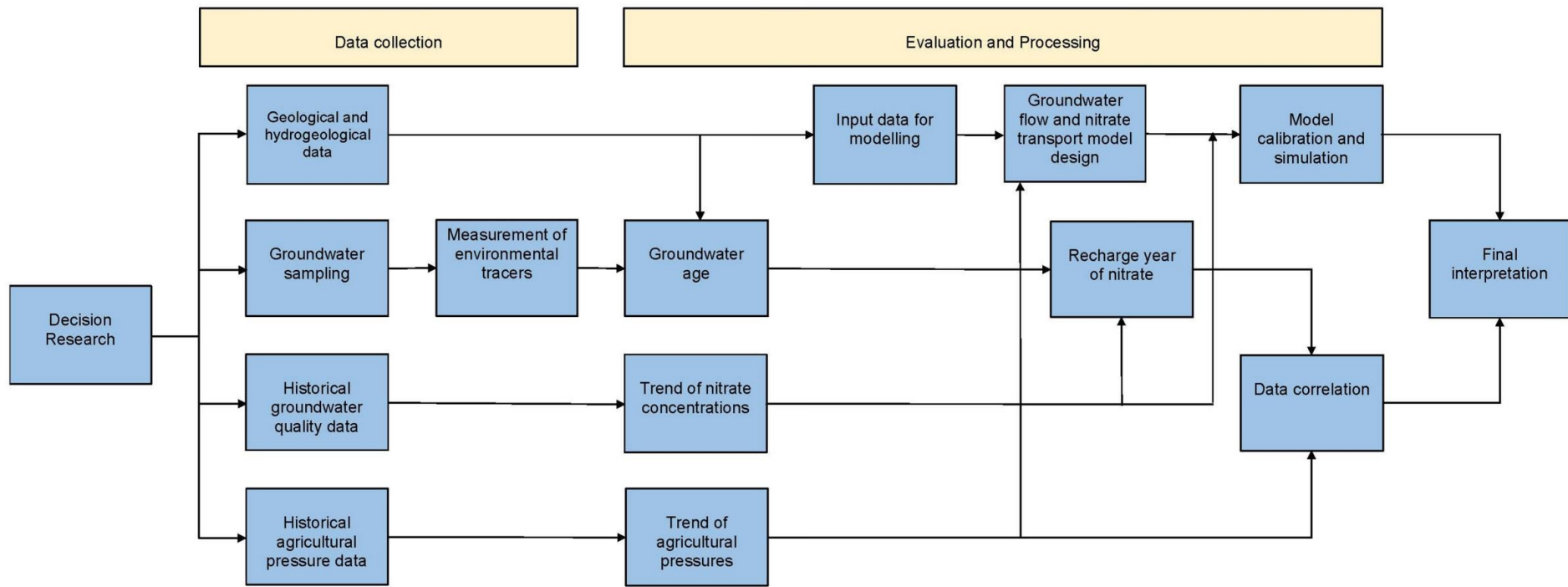


Figure 4

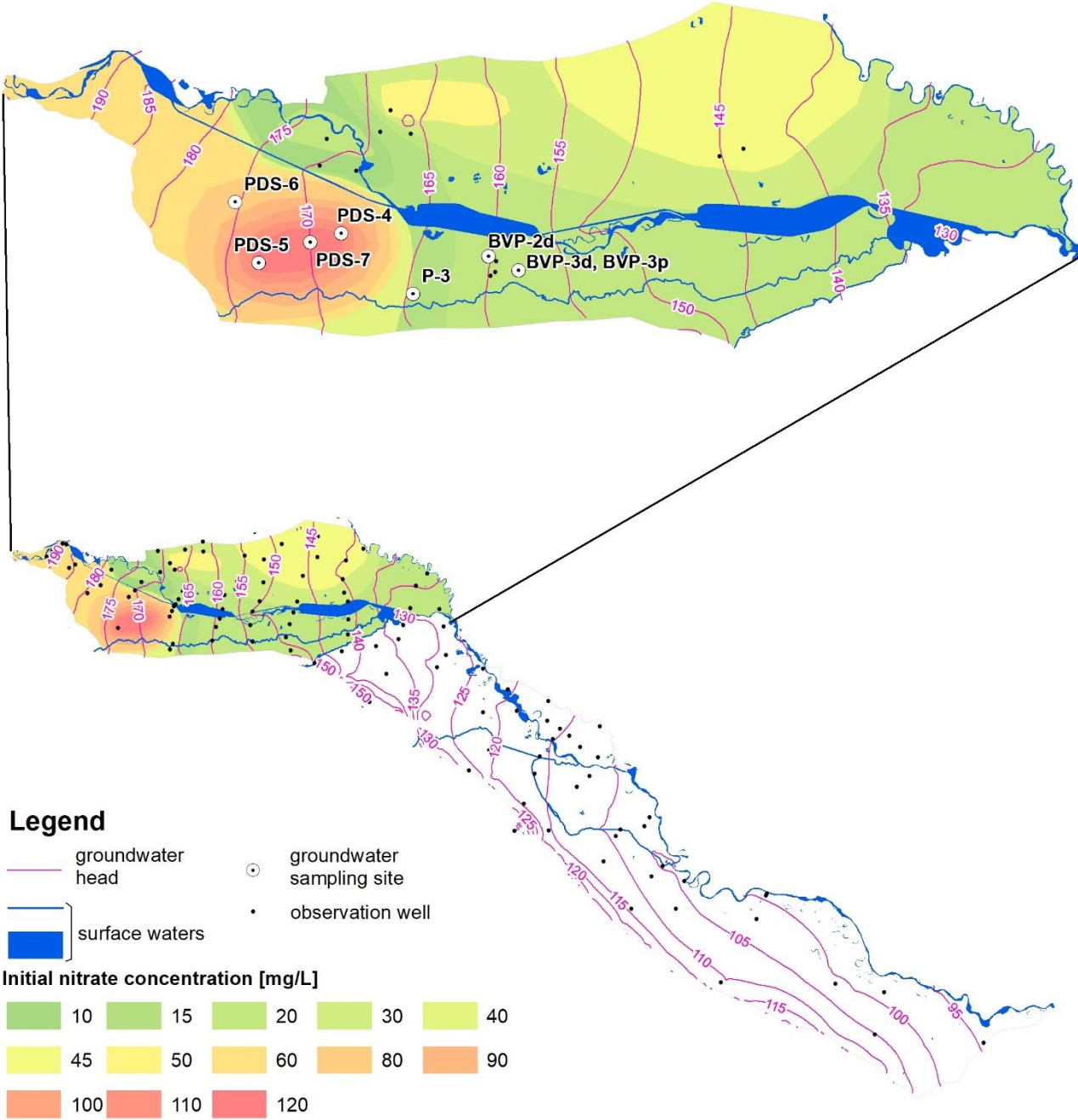


Figure 5

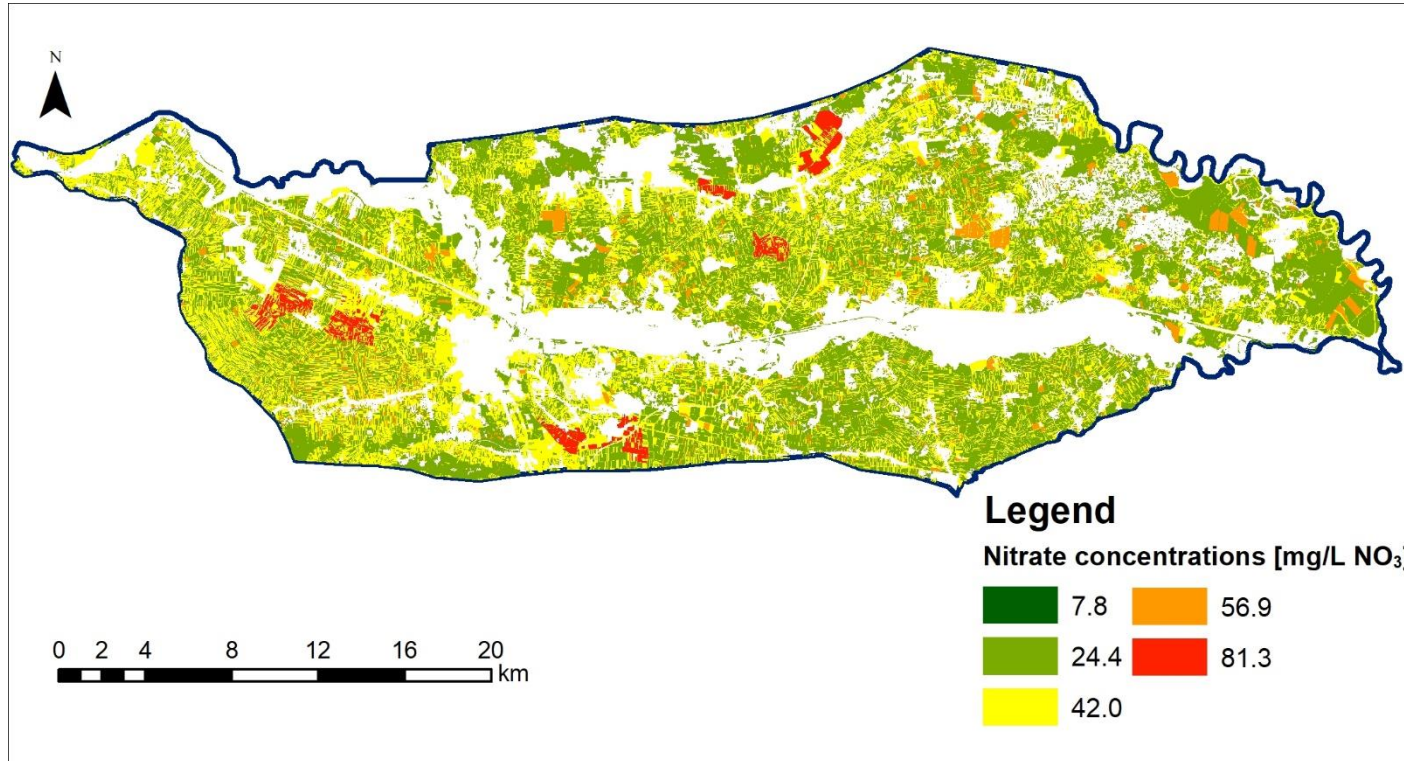
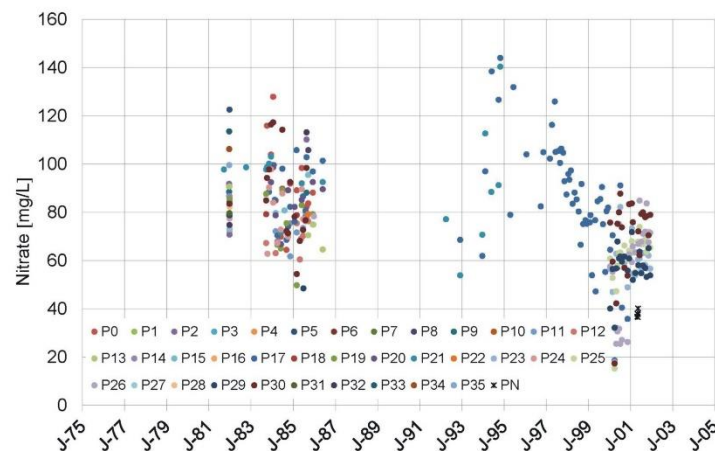
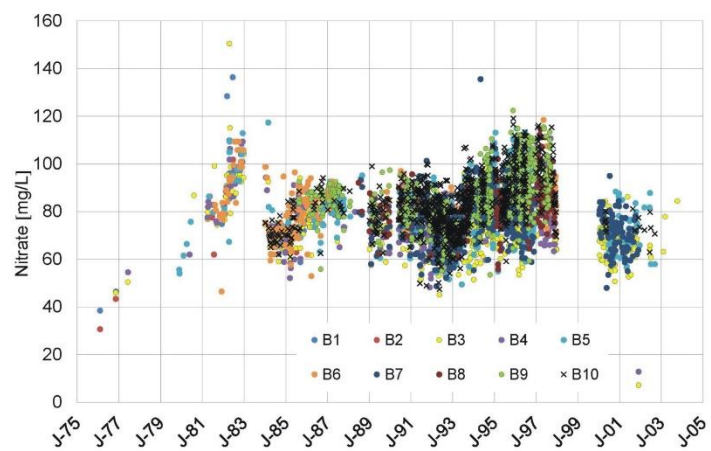


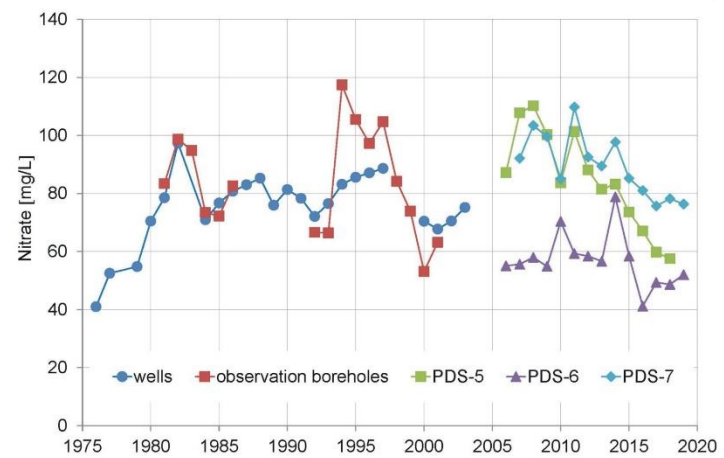


Figure 6

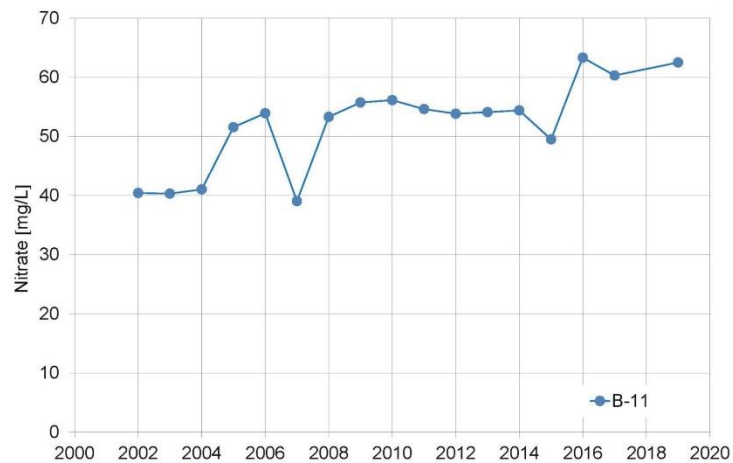


(a)

(b)

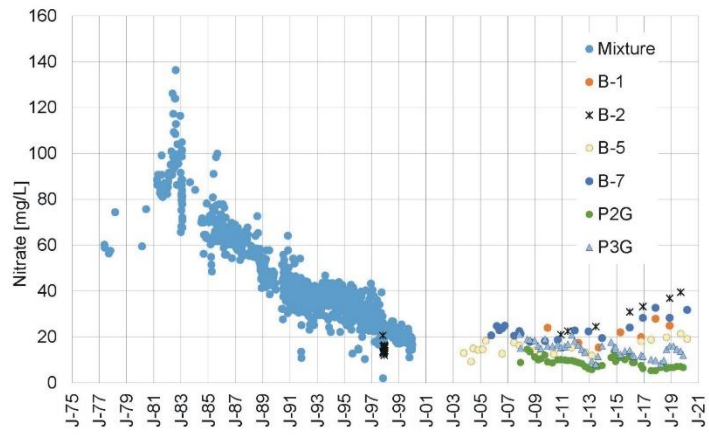


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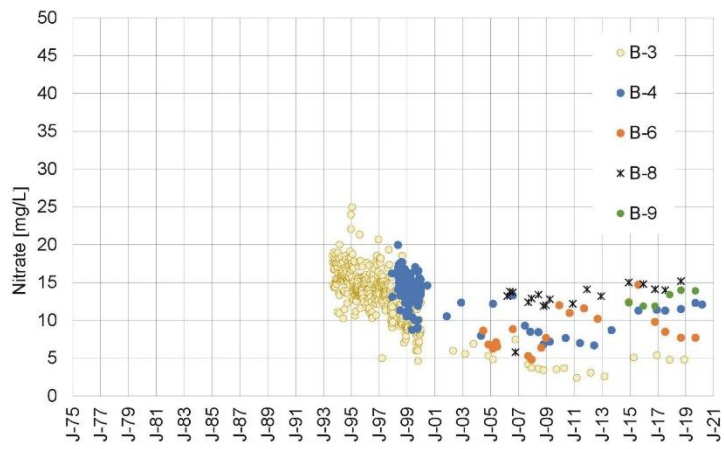


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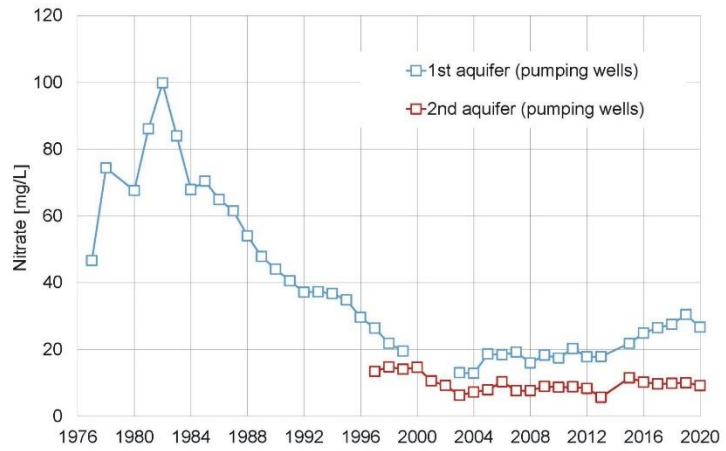
Figure 7



(a)

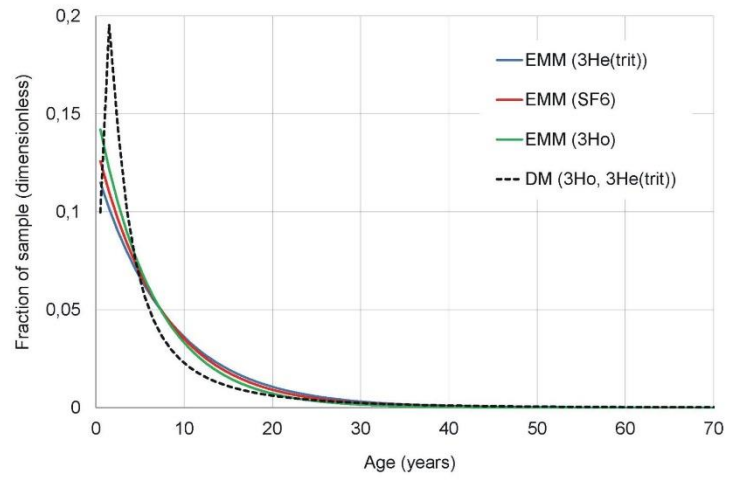


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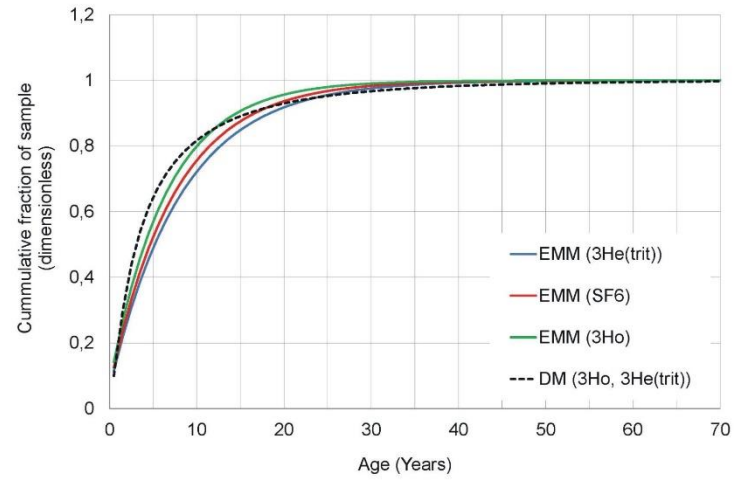


(c)

Figure 8

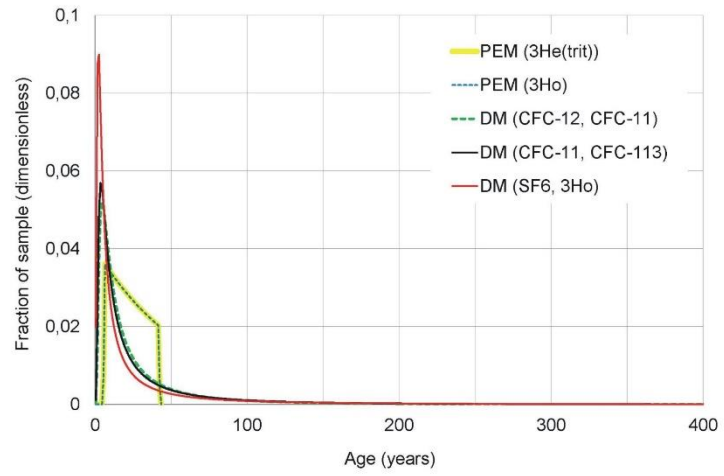


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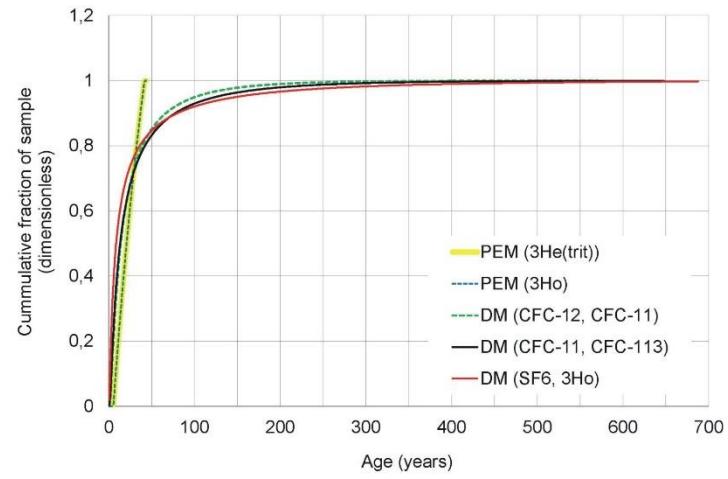


(b)

Figure 9

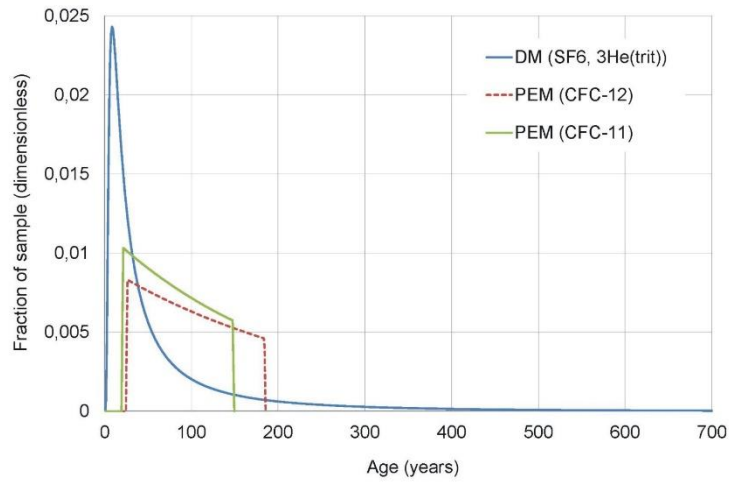


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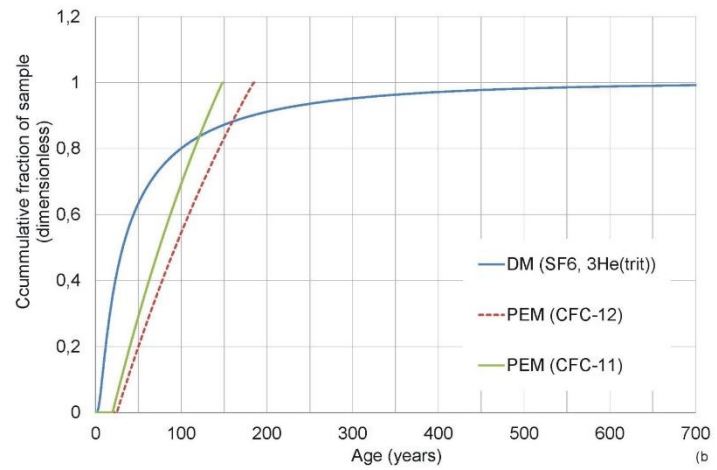


(b)

Figure 10

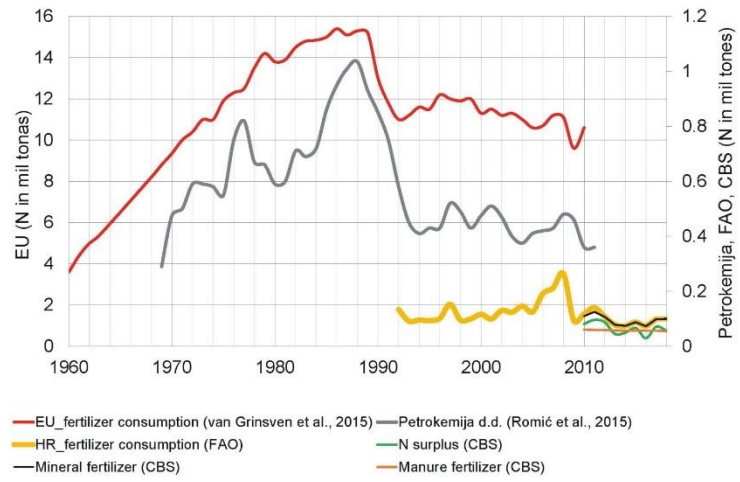


(a)

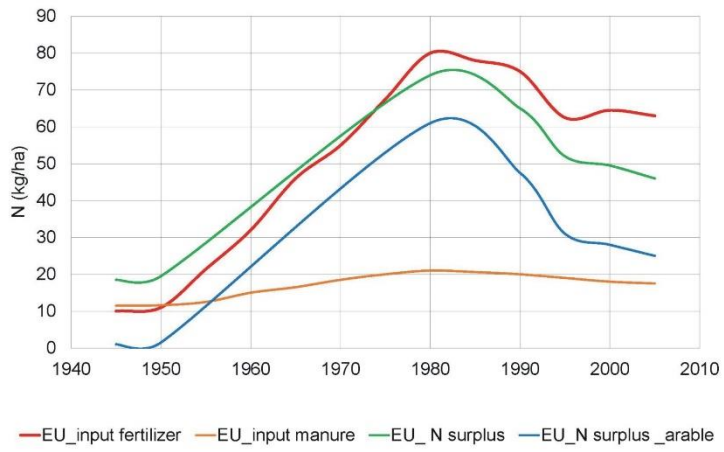


(b)

Figure 11



(a)



(b)

Figure 12

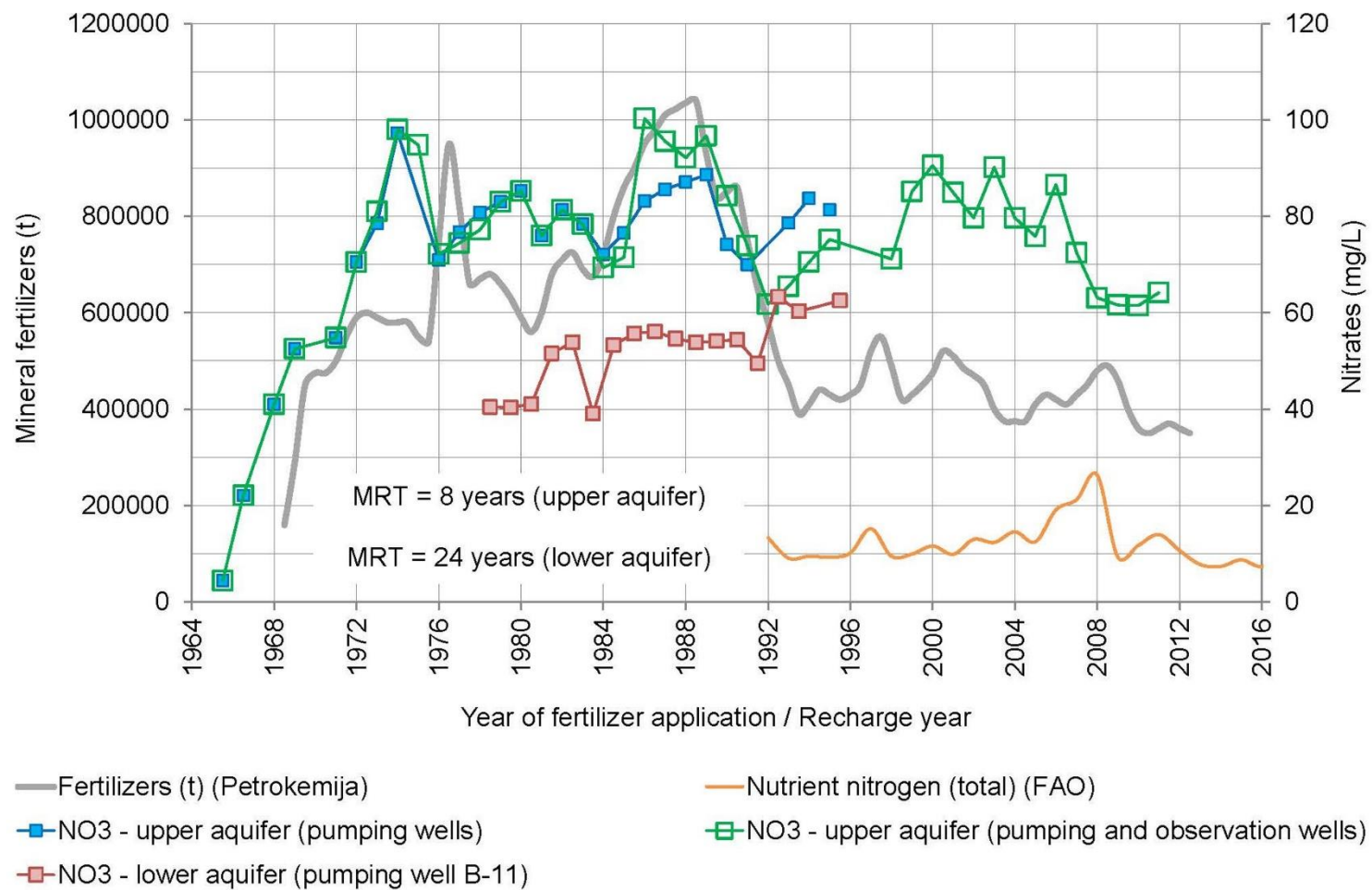


Figure 13

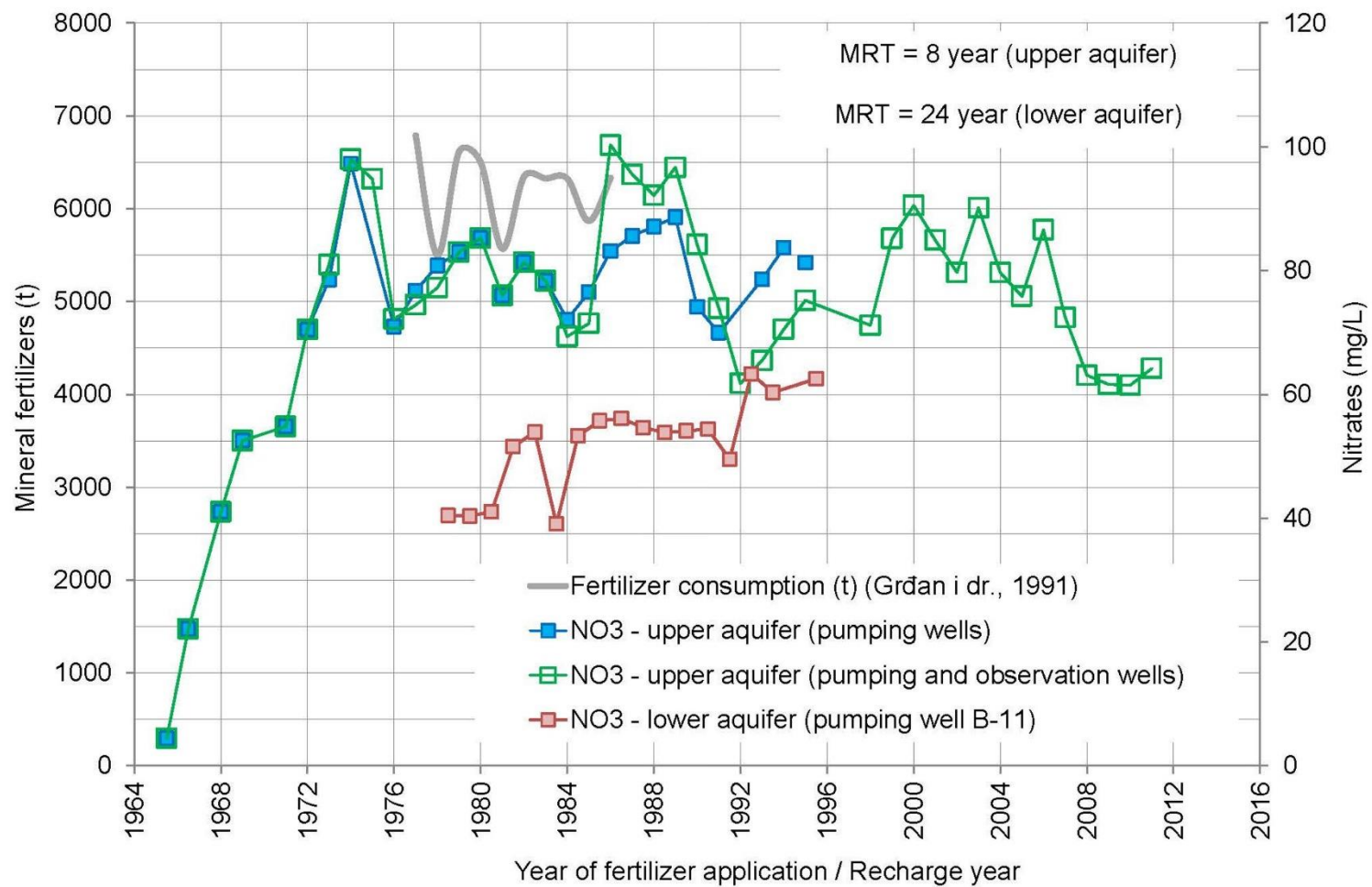




Figure 14

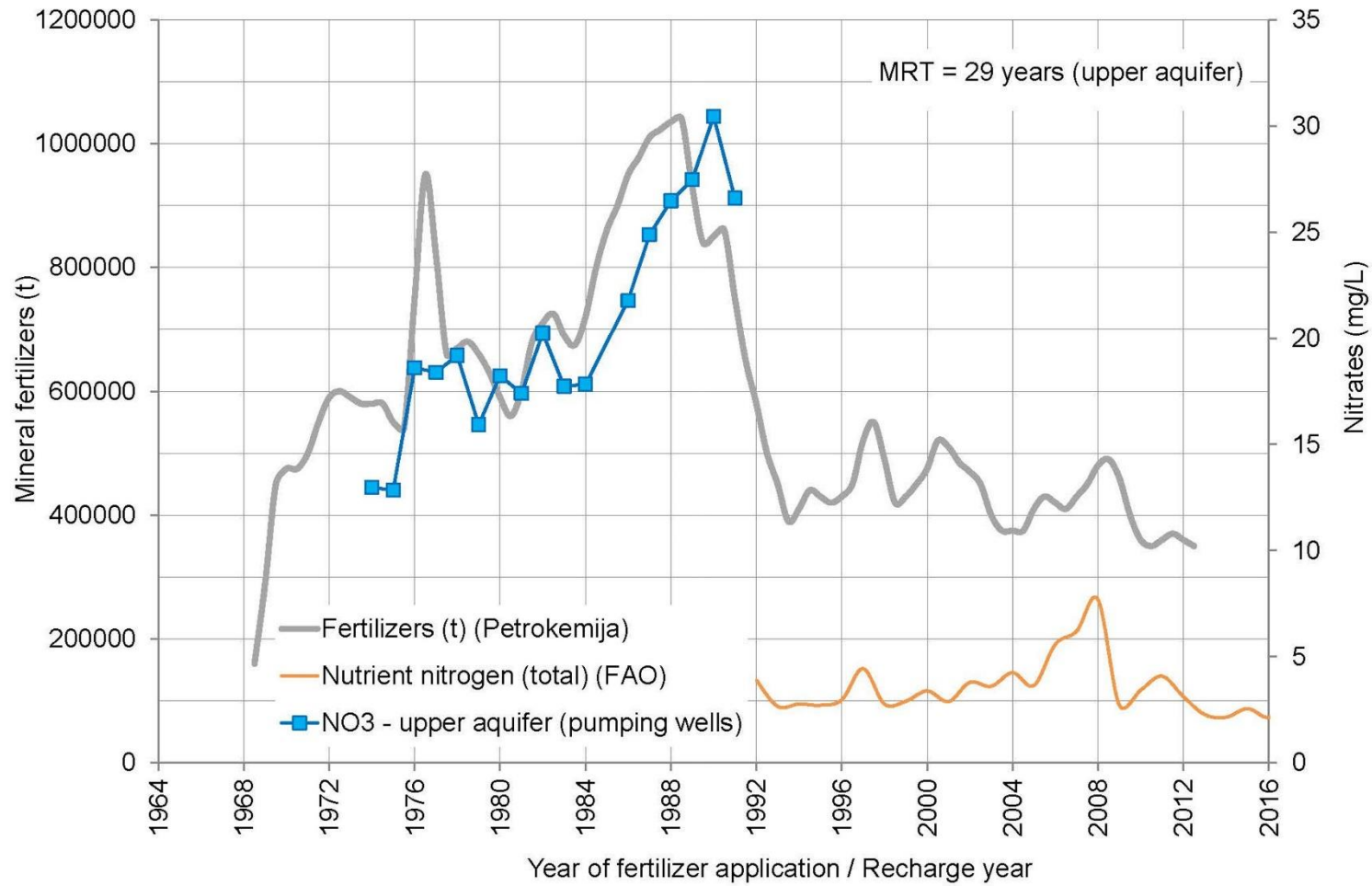
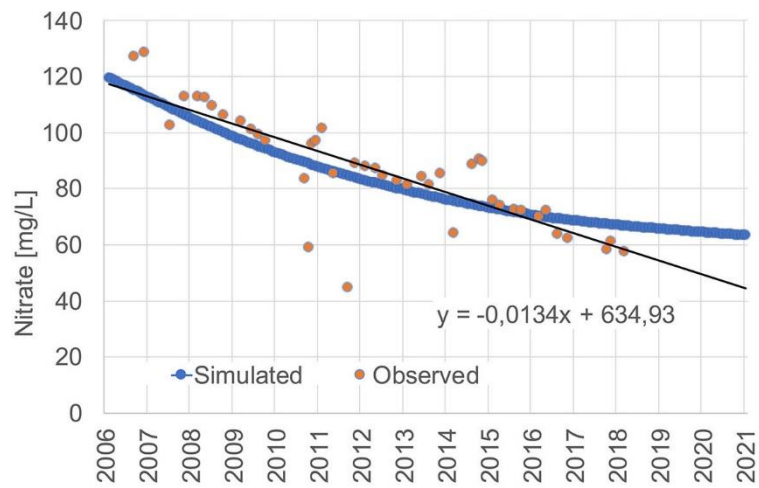
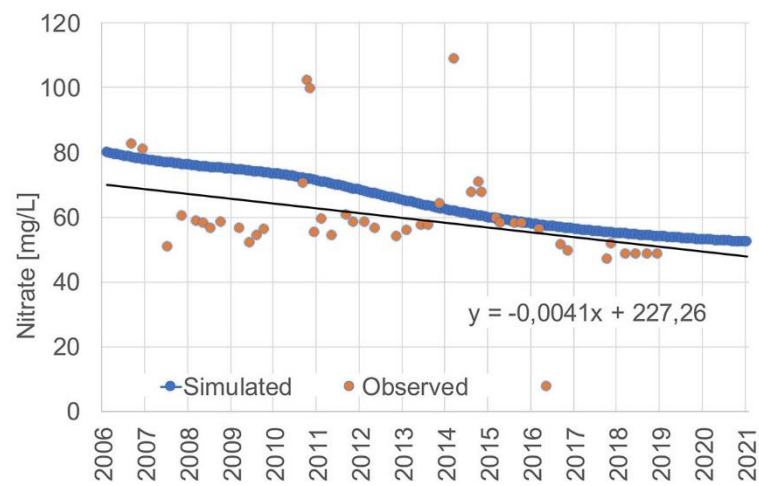


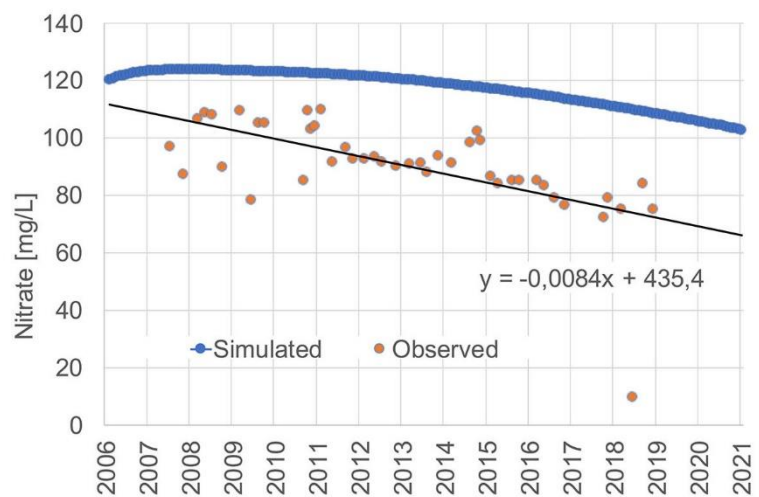
Figure 15



a)

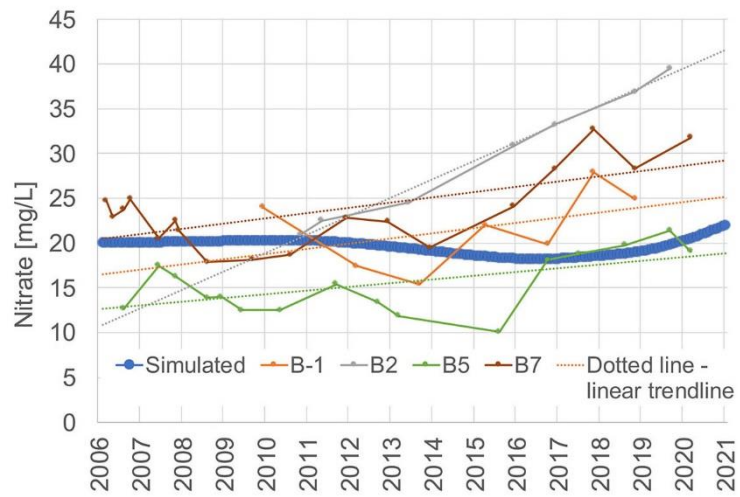


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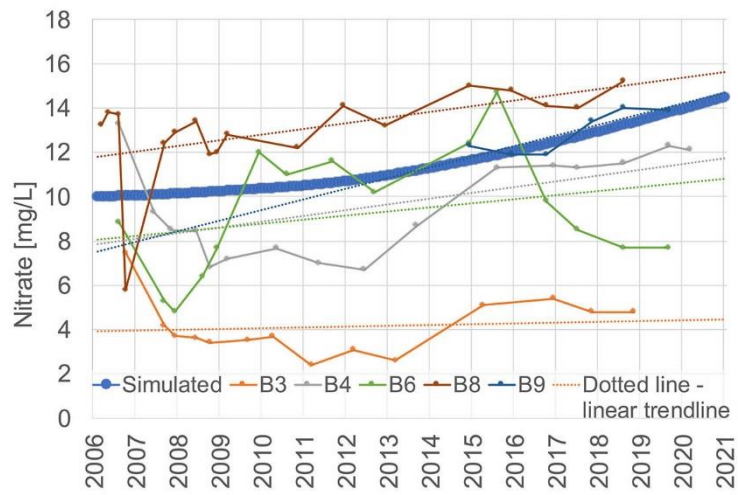


c)

Figure 16



a)



b)

Table 1. Characteristics of groundwater sampling sites

Observation borehole	Recharge elevation (m a.s.l)	Annual mean air temperature* (°C) 1960-1991/1971-2000	Depth of unsaturated zone (m)	Sampling date ( <sup>3</sup> H/ <sup>3</sup> He and CFCs)	Borehole depth (m)	Screen interval	Sampling depth (m)
PDS-5				11, 2017 and 03, 2018	26.0	13.7-19.7	15
PDS-6				11, 2017 and 03, 2018	24.0	11.7-17.7	15
PDS-7				11, 2017 and 03, 2018	34.0	29.3-32.3	15
PDS-4	170-200	9.9/10.2	2-6	11, 2019 and 03, 2020	66.7	42.0-46.0, 48.7-60.7	58
P-3D				11, 2019 and 03, 2020	40.0	4.0-19.0, 22.0-37.0	35
BVP-2D				11, 2019 and 03, 2020	71.0	57.0-69.0	50
BVP-3P				11, 2019 and 03, 2020	40.0	27.0-39.0	30
BVP-3D				11, 2019 and 03, 2020	70.0	58.0-70.0	58

\* Zaninović et al. (2008)

Table 2. Nitrate concentration data sets used to analyze the impact of agricultural pressures on groundwater

Observation borehole	Data time series	Data source
PDS-5		
PDS-6	2006-2019	Hrvatske vode (a legal entity for water management in Croatia)
PDS-7		
Wells and observation boreholes at the Varaždin and Bartolovec pumping sites	1973-2019	Varkom Ltd. for water supply and wastewater drainage (Varaždin)

Table 3. Agricultural lands and nitrate concentrations

Crop	Agricultural land use classes	Nitrate concentration in infiltrating water [mg/L NO <sub>3</sub> ]
Tobacco	Non-intensive land use	7.8
Cereals, corn, sugar beet, soy, oilseeds, potato, vineyards, meadows, pastures, sunflower	Extensive land use	24.4
Mosaic	A combination of extensive and intensive land use	42.0
Vegetables, cabbage, orchards	Intensive land use	56.9
Agricultural land in vicinity of farms	Nitrate on farmland	81.3

Table 4 Tracer concentrations in groundwater

Sample name	CFC conc. (pmol/L)			SF <sub>6</sub> (fmol/L)	<sup>3</sup> H (TU)	<sup>3</sup> H error (TU)	<sup>3</sup> H+ <sup>3</sup> He <sub>trit</sub> (TU)	<sup>3</sup> H+ <sup>3</sup> He <sub>trit</sub> error (TU)
	CFC-12	CFC-11	CFC-113					
PDS-5	<b>3.3 ± 0.2</b>	<b>10 ± 2</b>	0.42 ± 0.05	3.3 ± 0.4	4.91	0.09	10.6	0.9
PDS-6	<b>4.1 ± 0.3</b>	<b>19 ± 4</b>	0.43 ± 0.05	3.3 ± 0.4	5.05	0.09	7.9	0.8
PDS-7	<b>3.2 ± 0.2</b>	<b>10 ± 2</b>	0.43 ± 0.05	3.0 ± 0.3				
PDS-4	<b>3.3 ± 0.2</b>	<b>13 ± 3</b>	0.46 ± 0.05	1.6 ± 0.2	4.06	0.08	23.1	1.2
P-3D	<b>10 ± 3</b>	<b>30 ± 10</b>	0.34 ± 0.05	2.6 ± 0.3	4.90	0.18	27.2	1.4
BVP-2D	0.65 ± 0.05	0.8 ± 0.1	0.15 ± 0.05	2.2 ± 0.3				
BVP-3P	2.1 ± 0.2	5.3 ± 0.6	0.23 ± 0.05	3.4 ± 0.4	4.43	0.18	17.3	1.1
BVP-3D	0.42 ± 0.05	1.3 ± 0.2	0.04 ± 0.05	1.3 ± 0.2	1.63	0.06	53.5	2.0

Table 5. Calculated travel time range for groundwater samples and comparison with the calculated MRT using the environmental tracers

Sample name	Sampling depth below water table (m)	Recharge rate (mm/year)	Porosity (dimensionless)	Travel time (years)	MRT defined by the environmental tracers (years)
PDS-5	12	200-300	0.23	9.2 - 13.8	(29)* 6.5 - 13.7
PDS-6	9			6.9 - 10.4	(29)* 6.5 - 8.0
PDS-7	10			7.7 - 11.5	(29)* 9.0
PDS-4	49			37.6 - 56.4	24.5 - 30.9
P-3D	33			25.3 - 38.0	(14.5)** 30.4 - 33.0
BVP-2D	44			33.7 - 50.6	(18.5)** 39.5 - 53.0
BVP-3P	27			20.7 - 31.1	(7.5)** 24.2 - 36.0
BVP-3D	53			40.6 - 61.0	(27.5)** 49 - 62.1

\* (MRT defined using CFC-113)

\*\* (MRT defined using SF<sub>6</sub>)



Table 6 Summary of modelled MRTs extracted from different tracer combinations. PM: piston flow model; EMM: exponential mixing model; PEM: partial exponential model; DM: dispersion model.

Sample name	LPM	MRT		Modeled tracer	Mean value of MRT (years)
		(years)	Error (years)		
PDS-5	EMM	8.2	0.5	$^3\text{He}(\text{trit})$	8
PDS-5	EMM	6.5	1.3	$^3\text{H}_0$	
PDS-5	DM	7.0	0.03	$^3\text{H}_0, ^3\text{He}(\text{trit})$	
PDS-6	EMM	6.8	1.1	$^3\text{He}(\text{trit})$	
PDS-6	DM	9.4	0.03	$^3\text{H}_0, ^3\text{He}(\text{trit})$	
BVP-3P	PEM	22.2	0.4	$^3\text{He}(\text{trit})$	29
BVP-3P	PEM	22.2	0.6	$^3\text{H}_0$	
BVP-3P	DM	28.4	0.3	CFC-12, CFC-11	
BVP-3P	DM	32.3	0.3	CFC-11, CFC-113	
BVP-3P	DM	34.3	0.0	$\text{SF}_6, ^3\text{H}_0$	
BVP-2D	DM	69.7	0.3	CFC-12, CFC-113	
PDS-4	PEM	25.3	2.8	$\text{SF}_6$	24
PDS-4	PEM	24.0	0.3	$^3\text{He}(\text{trit})$	
PDS-4	PEM	23.9	0.5	$^3\text{H}_0$	
BVP-3D	PEM	78.1	6.8	CFC-11	84
BVP-3D	PEM	97.2	4.9	CFC-12	
BVP-3D	DM	75.6	9.5	$\text{SF}_6, ^3\text{He}(\text{trit})$	